TECHNICAL REPORT STANDARD PAGE 1. Report No. 3. Recipient's Catalog No. 2. Government Accession No. FHWA/LA-97/300 4. Title and Subtitle 5. Report Date ESTIMATION OF FREQUENCY BASED FLOOD PEAKS FOR June 1997 AN UNGAUGED WATERSHED USING FIELD 6. Performing Organization Code CALIBRATION 7. Author(s) 8. Performing Organization Report No. Menglou Wang and Fang Xin Yu 300 9. Performing Organization Name and Address 10. Work Unit No. Louisiana Transportation Research Center 4101 Gourrier Avenue 11. Contract or Grant No. Baton Rouge, LA 70808 92-1IMP 12. Sponsoring Agency Name and Address 13. Type of Report and Period Covered Louisiana Department of Transportation and Development **Final Report** Post Office Box 94245 July 1993-March 1997 Baton Rouge, LA 70804-9245 14. Sponsoring Agency Code 15. Supplementary Notes Conducted in cooperation with the U.S. Department of Transportation, Federal Highway Administration. 16. Abstract The present study has been conducted to evaluate eight flood prediction models for an ungauged small watershed. These models are either frequently used by or were developed by Louisiana Department of Transportation and Development (LADOTD). The eight models were applied to calculate flood frequency for the watershed on Ward Creek at Government Street in Baton Rouge. By comparing the results of the models with the flood peaks derived using the systematic flood records observed at the Ward Creek gauge station and by following the U.S. Water Resources Council (WRC) procedure, it was found that four of the eight models have better accuracy than the others. These four models are the U.S. Geological Survey (USGS) seven-parameter model, the USGS three-parameter model, the Lowe model and the Neely model. The U.S. Soil Conservation Service (SCS) model is widely used for flood prediction for small urban watersheds, but the accuracy of the model for the Ward Creek watershed is relatively low. This study shows that the accuracy of the SCS model can be significantly improved with the parameters calibrated using short-term field data. A procedure of parameter calibration using one- to two-year field data was developed in this study. The procedure may be used for more accurate flood prediction for watersheds with short-term stream gauging data or watersheds with long-term stream gauging data that have undergone significant hydrological changes. 17. Key Words 18. Distribution Statement rainfall-runoff, flood frequency, flood peaks, surface Unrestricted. This document is available through the runoff, small watersheds, Louisiana National Technical Information Service, Springfield, VA 21161. 19. Security Classif. (of this report) 20. Security Classif. (of this page) 21. No. of Pages 22. Price

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# ESTIMATION OF FREQUENCY BASED FLOOD PEAKS FOR AN UNGAUGED WATERSHED USING FIELD CALIBRATION

FINAL REPORT

by

MENGLOU WANG, P.E.
RESEARCH ASSOCIATE
LOUISIANA TRANSPORTATION RESEARCH CENTER

and

FANG XIN YU, PH.D., P.E., P.H.
RESEARCH ASSOCIATE
LOUISIANA TRANSPORTATION RESEARCH CENTER

LTRC REPORT NUMBER 300 LTRC PROJECT NUMBER 92-1 IMP STATE PROJECT NUMBER 736-99-0414 LSU PROJECT NUMBER 127-99-4174

conducted by

LOUISIANA DEPARTMENT OF TRANSPORTATION AND DEVELOPMENT LOUISIANA TRANSPORTATION RESEARCH CENTER

in cooperation with

U.S. DEPARTMENT OF TRANSPORTATION FEDERAL HIGHWAY ADMINISTRATION

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June 1997

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### **ABSTRACT**

The present study has been conducted to evaluate eight flood prediction models for an ungauged small watershed. These models are either frequently used by or were developed by Louisiana Department of Transportation and Development (LADOTD). The eight models were applied to calculate flood frequency for the watershed on Ward Creek at Government Street in Baton Rouge. By comparing the results of the models with the flood peaks derived using the systematic flood records observed at the Ward Creek gauge station using the U.S. Water Resources Council (WRC) procedure, it was found that four of the eight models have better accuracy than the others. These four models are the U.S. Geological Survey (USGS) seven-parameter model, the USGS three-parameter model, the Lowe model and the Neely model. The U.S. Soil Conservation Service (SCS) model is widely used for flood prediction for small urban watersheds, but the accuracy of the model for the Ward Creek watershed is relatively low. This study shows that the accuracy of the SCS model can be significantly improved with the parameters calibrated using shortterm field data. A procedure of parameter calibration using one- to two-year field data was developed in this study. The procedure may be used for more accurate flood prediction for watersheds with short-term stream gauging data or watersheds with long-term stream gauging data that have undergone significant hydrological changes.

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### **ACKNOWLEDGMENTS**

Funding for this project was supported by the Louisiana Transportation Research Center (LTRC) under LTRC project number 92-1 IMP. The Project Review Committee (PRC) consists of Henry J. Barousse, Zahir "Bo" Bolourchi, and Jack Manno. Stage and discharge data were provided by the USGS Baton Rouge District Office. Curtis Fletcher, Mark Morvant and David Bates at LTRC provided valuable support and guidance throughout the research. This project was started in July 1993 with Babak Naghavi serving as the original principal investigator. From December 1993 to December 1994, Fang Xin Yu was the principal investigator. Since May 1995, Menglou Wang has been the principal investigator. Babak Naghavi and Josh Gilbert (USGS Baton Rouge District Office) reviewed this report and provided valuable comments. The combined efforts of these individuals contributed greatly toward the successful completion of this project.

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### **IMPLEMENTATION STATEMENT**

The results of this study may assist LADOTD design engineers in selecting a flood prediction model for highway drainage design. The procedure of flood prediction developed in this study for ungauged watersheds with parameters calibrated using short-term field data demonstrated more accuracy and reliability than other methods that did not calibrate parameters. The improvement of the accuracy of flood frequency prediction is especially significant for models developed without using local field data, such as the SCS model.



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# CHAPTER 1 INTRODUCTION

Accurate estimations of flood peak, frequency, and volume are needed for safe and economical design of highway drainage and flood control structures. For watersheds with systematic stream gauging records of sufficient length, flood frequency analysis can be conducted by following the U.S. Water Resources Council (WRC) procedure [1]. The WRC procedure uses log-Pearson Type III distribution as a base method for flood frequency studies.

Accurate flood frequency analysis is more difficult for ungauged watersheds or watersheds with significant changes in land use or in drainage systems. Hydraulics design engineers often need to resort to information transfer techniques or regionalization procedures to estimate flood peaks and hydrographs. The information transfer techniques may consist of: (1) regional regression equations correlating peak discharge to climatic and watershed parameters [2], [3], [4], [5], (2) rainfall-runoff models calibrated by watershed characteristics [6], (3) regionalization models using data from nearby watersheds that exhibit similar climatic and hydrogeologic conditions [7], [8], and (4) the transfer of a flood frequency curve from nearby gauging stations. Each flood prediction model has its own assumptions and calibration conditions upon which the model was developed. Hence, estimated discharges from a watershed by different models may vary substantially. In many cases, designs may be over- or underestimated by 50 percent or more.

The accuracy of model prediction is heavily dependent on the accuracy of model parameter estimation. For an ungauged watershed, model parameter estimation is a rather difficult but necessary task in flood frequency analysis. Although tables and nomographs for parameter estimation are often provided by each model, determination of model parameters from these tables or nomographs is subject to large errors. Therefore, model calibration using short-term field data may be a better approach. If an accurate estimation of flood frequency is technically necessary and economically justified for a proposed

project, one-year or longer rainfall-runoff data may be observed to calibrate model parameters. This is especially true for the south central region of the United States where less variation of annual rainfall is experienced.

This report presents: (1) calculation of flood frequency for the Ward Creek watershed using eight flood prediction models, (2) establishment of the rating curve (stage-discharge relation) for the Ward Creek watershed, (3) evaluation of these flood prediction models, and (4) a procedure to apply a flood prediction model with parameters calibrated by using short-term field data. The eight selected models are (1) the Neely model [2], (2) the Lowe model [3], (3) the USGS seven-parameter model [4], (4) the USGS three-parameter model [4], (5) the Lee model [5], (6) the U.S. Soil Conservation Service (SCS) model [6], (7) the Louisiana Regional GEV-PWM model [7], and (8) the Louisiana GEV-OPT model [8].

A rural watershed on Beaver Bayou above Hooper Road in Baton Rouge was also selected at the beginning of the study. During the study, it was discovered that Beaver Bayou overbanks almost every year, so a proper rating curve cannot be established. The results for the Beaver Bayou watershed are not included in this report.

# CHAPTER 2 OBJECTIVES

The objectives of this study are:

- (1) To compute flood magnitudes at the return periods of 2, 5, 10, 25, 50, and 100 years for an ungauged watershed using eight watershed models. The eight models are:
  - (a) the Neely model [2]
  - (b) the Lowe model [3]
  - (c) the USGS seven-parameter model [4]
  - (d) the USGS three-parameter model [4]
  - (e) the Lee model [5]
  - (f) the SCS model [6]
  - (g) the Regional GEV-PWM model [7] and
  - (h) the Regional GEV-OPT model [8];
- (2) To evaluate the eight models by comparing the model results with those derived from the WRC procedures [1] using long-term stream gauging data. The advantages and disadvantages of each model will be discussed; and
- (3) To develop a field-calibration procedure for flood prediction when more accurate estimation of flood peaks are economically justified. This requires the designer(s) to set up a network of rain gauges and a flowmeter to collect rainfall-runoff data for one year or more. The model parameters are calibrated using short-term rainfall-runoff data to improve the accuracy of the model.

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# CHAPTER 3 SCOPE

The scope of this project encompasses selection of eight frequently used flood prediction models, selection of a small watershed, installation of three rain gauges and a flowmeter on the watershed, computation of flood magnitudes and frequencies for the watershed using the eight models, evaluation of the eight models by comparing the model results with the flood peaks derived using the WRC procedure, and development of a procedure to improve flood prediction accuracy by calibrating model parameters using short-term field data.

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# CHAPTER 4 DATA COLLECTION

#### Watershed Selection

The Ward Creek watershed is an urban watershed located in downtown Baton Rouge (Figure 4.1). The outlet is on the main channel of Ward Creek at Government Street. The drainage area is 4.63 square miles (12.0 km²). In 1968 the main channel of Ward Creek was upgraded from a natural ditch to a concrete lined channel. As shown in Figure 4.2, the cross-section of the lined main channel at the flowmeter installation site 200 ft (61 meters) upstream from Government Street is trapezoidal with a bottom width of 24.1 ft (7.3 m), top width of 62.0 ft (18.9 m), and depth of 11.4 ft (3.5 m). The USGS has maintained a gauge station on Ward Creek at Government Street since 1954, with a station number of 7379000.

## **Rainfall Data Collection**

Three rain gauges (ISCO 674L Model) were installed on the watershed. A picture of the rain gauge is shown in Appendix 1(a). The locations of the rain gauges are shown in Figure 4.1, and they are listed below:

Rain Gauge A: 265 S. Foster Drive (State Police Headquarters)

Rain Gauge B: 2654 Mission Drive (Magnolia Construction Company)

Rain Gauge C: 4045 Scenic Highway (Exxon Baton Rouge Refinery)

From February 1994 to July 1995, hourly rainfall data were recorded by the rain gauges. From July 1995 to December 1996, the rainfall recording interval was 10 minutes.

## Stage and Discharge Data Collection

Stage and discharge data for the Ward Creek gauge station were obtained from the USGS Baton Rouge District Office. From 1954 to 1967, the USGS recorded continuous stage and discharge data for the Ward Creek watershed. In 1968, there was a one-year data gap during the channel lining construction. Since 1969, only continuous stage data have been recorded by USGS.

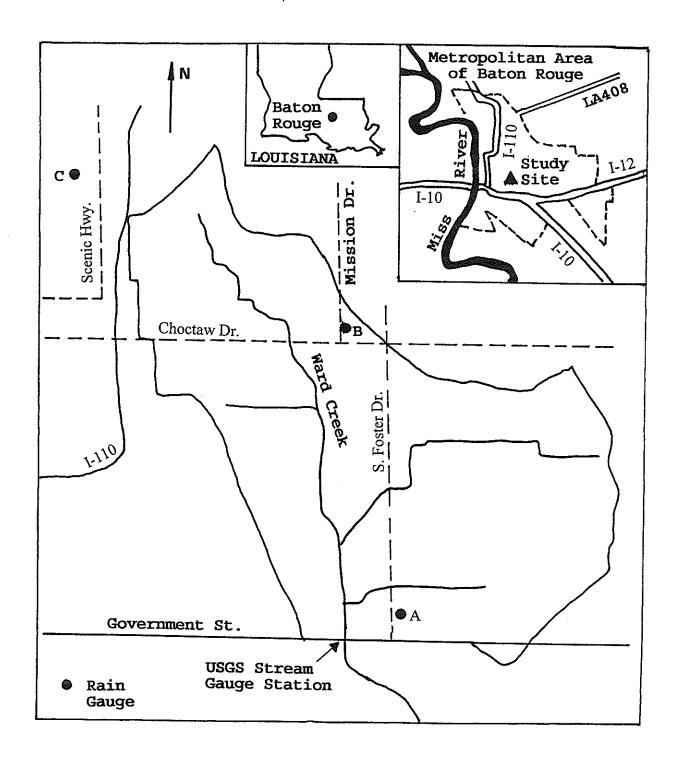


Figure 4.1
The Ward Creek watershed above Government Street in Baton Rouge

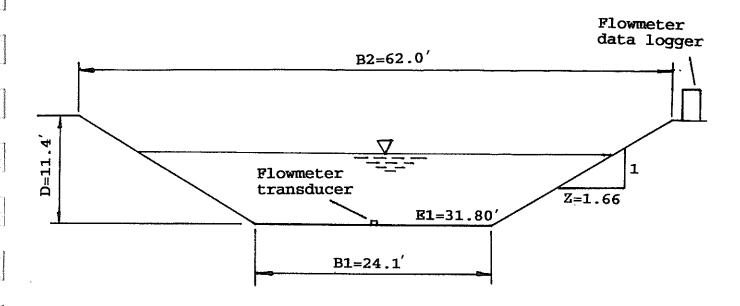


Figure 4.2
Cross-section of Ward Creek at Government Street

In order to convert the USGS stage data from 1969 through 1995 to discharge data, a flowmeter with data logger (American Sigma 950AV Model) was purchased to establish the rating curve. A picture of the flowmeter is shown in Appendix 1(b). The flowmeter is capable of measuring both water level and flow velocity. As described in the manufacturer's manual, the Doppler ultrasonic transducer measures the average flow velocity near the transducer. The error of velocity measurements is  $\pm 2$  percent and the error of water level measurements is  $\pm 0.012$  ft. The rating curve was established based on the water level and flow velocity data collected from May to August 1996.

#### Other Data

The geometric parameters of the watersheds (drainage area, channel length, and main channel slope) were obtained from topographic maps published by USGS in cooperation with LADOTD. The land use data for the two watersheds were obtained from East Baton Rouge Parish Department of Public Works.

#### **CHAPTER 5**

## SELECTED FLOOD PREDICTION MODELS

Eight flood estimation models frequently used by LADOTD or developed with the involvement of LADOTD are selected to estimate flood peaks for the Ward Creek watershed. These models were developed for ungauged small watersheds. Each model is briefly described in this chapter. The reader should refer to the original reports or papers for more details.

## 1. The Neely Model

Neely [2] developed a model to estimate the magnitudes of floods for streams in Louisiana with recurrence intervals of 2 to 100 years. Data from 170 gauging stations in Louisiana, with drainage areas between 0.01 and 3,000 square miles (0.026 and 7,800 square kilometers), were used in the analysis. Flood magnitudes for different return periods at each gauging station were computed by fitting a log-Pearson Type III distribution. The model for unregulated or rural watersheds was developed by multiple regression technique.

$$Q_2 = 3.40A^{0.72}(P-35)^{0.95}S^{0.38}$$
 (5-1a)

$$Q_5 = 4.77A^{0.76}(P-35)^{0.88}S^{0.54}$$
 (5-1b)

$$Q_{10} = 5.04 A^{0.79} (P - 35)^{0.88} S^{0.63}$$
 (5-1c)

$$Q_{25} = 5.31 A^{0.81} (P-35)^{0.89} S^{0.71}$$
 (5-1d)

$$Q_{50} = 5.40 A^{0.82} (P - 35)^{0.90} S^{0.75}$$
 (5-1e)

$$Q_{100} = 5.32 A^{0.83} (P-35)^{0.92} S^{0.80}$$
 (5-1f)

where  $Q_x$  = peak discharge (cfs) for a corresponding rural watershed for a recurrence interval of x years

A = drainage area (square miles)

P = mean annual precipitation (inches)

S = main channel slope (ft/mile).

The main channel slope is measured between two points along the main channel from the outlet to the upper divide, one point at 10 percent of the channel length and the other at 85 percent of the channel length. When the main channel slope is less than 0.5 ft/mile, 0.5 is used. When the main channel is greater than 30 ft/mile, 30 is adopted. For urban watersheds, Neely [2] modified Sauer's [9] urban model and recommended the following flood estimation equations:

$$Q_{2u} = R_L Q_2 \tag{5-2a}$$

$$Q_{5u} = 1.66(R_L - 1)Q_2 + 0.167(7 - R_L)Q_5$$
 (5-2b)

$$Q_{10\mu} = 1.97(R_L - 1)Q_2 + 0.167(7 - R_L)Q_{10}$$
 (5-2c)

$$Q_{25u} = 2.38(R_L - 1)Q_2 + 0.167(7 - R_L)Q_{25}$$
 (5-2d)

$$Q_{50u} = 2.68(R_L - 1)Q_2 + 0.167(7 - R_L)Q_{50}$$
 (5-2e)

$$Q_{100u} = 2.98(R_L - 1)Q_2 + 0.167(7 - R_L)Q_{100}$$
 (5-2f)

where  $Q_{xu}$  = peak discharge (cfs) for an urban watershed at the return period of x years;  $Q_x$  = peak discharge (cfs) for a corresponding rural watershed at the return period of x years calculated using a rural runoff model such as Equation (5-1a) through Equation (5-1f);  $R_L$  = urban adjustment ratio, which was defined as the ratio of mean annual urban peak discharge,  $Q_{2.33u}$ , to the mean annual natural peak discharge,  $Q_{2.33}$  [10]. The percentage of impervious area in the basin and the percentage of the basin served by storm sewers (including improved channels) are needed to estimated  $R_L$ . The urban adjustment ratio can be determined from Figure 15 in Neely's report [2].

### 2. The Lowe Model

Regression equations for estimating magnitude and frequency of peak discharge for small streams in Louisiana were developed by Lowe [3]. Three independent parameters were used for non-linear multiple parameter regression. The three parameters are the main channel slope, drainage area, and mean annual precipitation. The following regression equations were obtained by Lowe

[3] for watersheds that drain less than 10 square miles (25.9 km²), have main channel slopes between 5 and 100 ft/mile, and are not affected by regulation.

$$Q_2 = 5.50 A^{0.56A^{-0.08}} S^{1.10S^{-0.20}} (P-35)^{0.69}$$
 (5-3a)

$$Q_5 = 8.32 A^{0.61A^{-0.08}} S^{1.25S^{-0.20}} (P-35)^{0.63}$$
 (5-3b)

$$Q_{10} = 10.2A^{0.63A^{-0.06}}S^{1.28S^{-0.20}}(P-35)^{0.62}$$
 (5-3c)

$$Q_{25} = 12.9A^{0.64A^{-0.04}}S^{1.30S^{-0.20}}(P-35)^{0.63}$$
 (5-3d)

$$Q_{50} = 12.9 A^{0.65 A^{-0.06}} S^{1.33 S^{-0.20}} (P-35)^{0.66}$$
 (5-3e)

$$Q_{100} = 12.6A^{0.67A^{-0.06}}S^{1.36S^{-0.20}}(P-35)^{0.68}$$
(5-3f)

where A, S and P were defined as in Neely [2]. To determine the flood peaks for an urban watershed, the flood peaks for a corresponding rural watershed are calculated using the Lowe model. Then the flood peaks for the urban watershed are calculated using the Neely model defined by Equation (5-2a) through Equation (5-2f).

## 3. The USGS Seven-Parameter Model

Sauer et al. [4] analyzed 269 gauged urban watersheds in 56 cities and 31 states in the U.S. and developed a seven-parameter regression model and a three-parameter regression model for estimating flood peaks at the return periods of 2, 5, 10, 25, 50, and 100 years. One watershed in Baton Rouge was included in the study by Sauer et al. [4]. The seven-parameter model developed by Sauer et al. [4] was defined by the following equations:

$$Q_{2u} = 2.35A^{0.41}S^{0.17}(I_2 + 3)^{2.04}(ST + 8)^{-0.65}(13-BDF)^{-0.32}IA^{0.15}Q_2^{0.47}$$
 (5-4a)

$$Q_{5u} = 2.70A^{0.35}S^{0.16}(I_2 + 3)^{1.86}(ST + 8)^{-0.59}(13-BDF)^{-0.31}IA^{0.11}Q_5^{-0.54}$$
 (5-4b)

$$Q_{10u} = 2.99A^{0.32}S^{0.15}(I_2 + 3)^{1.75}(ST + 8)^{-0.57}(13-BDF)^{-0.30}IA^{0.09}Q_{10}^{0.58}$$
 (5-4c)

$$Q_{25u} = 2.78A^{0.31}S^{0.15}(I_2 + 3)^{1.76}(ST + 8)^{-0.55}(13-BDF)^{-0.29}IA^{0.07}Q_{25}^{0.60}$$
 (5-4d)

$$Q_{50u} = 2.67A^{0.29}S^{0.15}(I_2 + 3)^{1.74}(ST + 8)^{-0.53}(13-BDF)^{-0.28}IA^{0.06}Q_{50}^{-0.62}$$
 (5-4e)

$$Q_{100u} = 2.50A^{0.29}S^{0.15}(I_2 + 3)^{1.76}(ST + 8)^{-0.52}(13-BDF)^{-0.28}IA^{0.06}Q_{100}^{0.63}$$
 (5-4f)

- where  $Q_{xu}$  = peak discharge (cfs) for an urban watershed at the return period of x years
  - A = drainage area (square miles)
  - S = main channel slope (ft/mile)
  - I<sub>2</sub> = rainfall intensity (inches) for the duration of 2 hours and return period of 2 years
  - ST = basin storage, the percentage of the watershed occupied by lakes, reservoirs, swamps, and wetlands
  - BDF = basin development factor, an index of the prevalence of the drainage aspects of (a) storm sewers, (b) channel improvements, (c)impervious channel linings, and (d) curb-and-gutter streets
    - IA = impervious area, percentage of the drainage basin occupied by impervious surface, such as houses, buildings, streets, and parking lots
    - $Q_x$  = peak discharge (cfs) for a corresponding rural drainage basin for the recurrence interval of x years.

The corresponding rural discharges are computed using other rural flood prediction models.

### 4. The USGS Three-Parameter Model

By dropping the less significant variables in the USGS seven-parameter model, Sauer *et al.* [4] developed a three-parameter model to estimate urban peak discharge.

$$Q_{2u} = 13.2A^{0.21}(13 - BDF)^{-0.43}Q_2^{0.73}$$
 (5-5a)

$$Q_{5u} = 10.6A^{0.17}(13 - BDF)^{-0.39}Q_5^{0.78}$$
 (5-5b)

$$Q_{10u} = 9.51 A^{0.16} (13 - BDF)^{-0.36} Q_{100}^{0.79}$$
 (5-5c)

$$Q_{25u} = 8.68A^{0.15}(13-BDF)^{-0.34}Q_{25}^{0.80}$$
 (5-5d)

$$Q_{50u} = 8.04A^{0.15}(13 - BDF)^{-0.32}Q_{50}^{0.81}$$
 (5-5e)

$$Q_{100u} = 7.70A^{0.15}(13 - BDF)^{-0.32}Q_{100}^{0.82}$$
 (5-5f)

where  $\Omega_{xu},\,A,\,BDF,$  and  $\Omega_{x}$  are defined the same as in the seven-parameter model.

### 5. The Lee Model

Lee [5] analyzed 217 stream-gauging sites in Louisiana and bordering states of Mississippi and Arkansas. Regression equations were developed using these watersheds that are basically natural and unregulated with drainage areas less than 3,000 square miles (7770 km²).

$$Q_2 = 5.45 A^{0.62} (P - 35)^{1.00} S^{0.33}$$
 (5-6a)

$$Q_5 = 5.50A^{0.68}(P-35)^{1.01}S^{0.51}$$
 (5-6b)

$$Q_{10} = 5.24 A^{0.71} (P - 35)^{1.03} S^{0.61}$$
 (5-6c)

$$Q_{25} = 4.85 A^{0.74} (P - 35)^{1.06} S^{0.71}$$
 (5-6d)

$$Q_{50} = 4.25A^{0.77}(P-35)^{1.10}S^{0.78}$$
 (5-6e)

$$Q_{100} = 3.85 A^{0.79} (P - 35)^{1.13} S^{0.84}$$
 (5-6f)

where  $Q_x$ , A, P and S are defined the same as in Neely [2] and Lowe [3]. With the results from Equations (5-6a) through (5-6f), the flood peaks for an urban watershed can be determined by using the Neely model defined by Equations (5-2a) through (5-2f).

#### 6. The SCS Model

The Soil Conservation Service Model [4], which was published in the SCS Technical Release 55 (TR-55), presents simplified procedures to calculate storm runoff volume, peak discharge, and hydrographs,. The SCS model was developed for urban small watersheds. The TR-55 report says that the model

applies to watersheds with time of concentration from 0.1 to 10 hours. Therefore, the SCS model should be applicable to the Ward Creek watershed because its time of concentration is about 3 hours (see Appendix 2).

The model starts with the assumption that a gross rain amount P (in inches) is uniformly distributed in the watershed. To account for the initial losses before runoff occurs and subsequent losses due to infiltration, depression, evapotranspiration, *etc.*, the amount of rainfall is converted to mass direct runoff R (in inches) or the excess rainfall by the following equation,

$$R = \frac{[P - 0.2(\frac{1000}{CN} - 10)]^2}{P + 0.8(\frac{1000}{CN} - 10)}$$
(5-7)

where CN is the runoff curve number, which depends on soil group, cover type, and antecedent moisture condition. CN can be determined by using Table 2.2, Figures 2.3 and 2.4 given in the SCS manual [6], and by using the local hydrological soil group maps obtainable from local SCS offices. Peak discharge can be estimated either by the graphical peak discharge method or by the tabular hydrograph method. If hydrograph is not of interest, peak discharge can be estimated by the graphical peak discharge method using the following equation:

$$Q_p = q_u A R F_p \tag{5-8}$$

where  $Q_p$  = peak discharge (cfs)

q<sub>u</sub> = unit peak discharge per square mile of drainage area per inch of runoff (cfs/mile².in), which can be estimated by using Exhibit 4-I,
 4-IA, 4-II or 4-III provided by the SCS manual [6]

A = drainage area (square miles)

R = direct runoff computed from Equation (5-7)

F<sub>p</sub> = pond and swamp adjustment factor, which can be determined from Table 4.2 in the SCS manual] [6].

## 7. The Louisiana Regional GEV-PWM Model

Naghavi *et al.* [7] divided Louisiana into four hydrologically homogeneous regions based on the analyses of soil, topographic, and climatic maps. The four regions are: southeast, southwest, northeast, and northwest. For each of the four regions, a regression equation was developed to estimate the mean annual maximum flood in terms of the drainage area.

For the southeast region,

$$Q_m = 10^{2.695A^{0.072}} (5-9a)$$

For the southwest region,

$$Q_m = 10^{2.561A^{0.076}}$$
 (5-9b)

For the northeast region,

$$Q_m = 10^{2.406A^{0.063}}$$
 (5-9c)

For the northwest region,

$$Q_m = 10^{2.836A^{0.062}}$$
 (5-9d)

where  $Q_m$  = mean annual maximum discharge (cfs)

A = drainage area (square miles).

The method is called GEV-PWM model because the regression coefficients were determined by the generalized extreme value (GEV) distribution with the probability weighted moments (PWM). With these four regression equations, the indexed regional optimization procedure developed in this study can be easily extended to predict flood frequency at ungauged sites. This can be done by using the following three steps:

Step 1: Estimate the drainage area, A, for the ungauged watershed and identify the region in which the watershed is located using Figure 3 in Naghavi et al. [7];

Step 2: Calculate the mean annual maximum discharge,  $Q_m$ , by using one of the equations from Equation (5-9a) to Equation (5-9d);

Step 3: Calculate the design discharge for a selected frequency by multiplying the regional flood quantile (dimensionless) at the same frequency by the mean annual maximum discharge calculated at Step 2. The dimensionless regional quantiles for the return periods of 2, 5, 10, 25, 50, and 100 years of flood for the four hydrologically homogeneous regions are given in Table 5.1.

This method did not consider urbanization. Both rural and urban watersheds were used for regression analysis of the GEV-PWM model.

## 8. The Louisiana Regional GEV-OPT Model

Naghavi and Yu [8] modified the Louisiana Regional GEV-PWM Model [7]. In the modified model, the dimensionless regional flood quantiles were estimated by using the GEV distribution and optimization method (GEV-OPT). Ninety gauged stream sites in Louisiana were selected for the study. To determine the peak discharge for return periods of 2, 5, 10, 25, 50, and 100 years for the four homogeneous regions in Louisiana, the three steps described previously in the Louisiana Regional GEV-PWM model are followed. Equation (5-9a) through Equation (5-9d) are used to calculate the mean annual maximum discharge. The dimensionless regional flood quantiles are listed in Table 5.2.

Table 5.1
Dimensionless regional flood quantiles for the GEV-PWM model

Return	Hydrologically homogeneous region				
period (years)	Southeast	Southwest	Northeast	Northwest	
2	0.811	0.758	0.978	0.712	
5	1.410	1.380	1.300	1.350	
10	1.920	1.910	1.470	1.965	
25	2.638	2.773	1.659	2.971	
50	3.237	3.584	1.778	3.957	
100	3.916	4.570	1.882	5.198	

Source: Naghavi et al. [7].

Table 5.2
Dimensionless regional flood quantiles for the GEV-OPT procedure

Return	Hydrologically homogeneous region				
period (years)	Southeast	Southwest	Northeast	Northwest	
2	0.812	0.763	0.999	0.708	
5	1.315	1.320	1.325	1.260	
10	1.932	1.980	1.544	2.018	
25	2.607	2.862	1.732	3.026	
50	3.162	3.674	1.845	3.989	
100	3.766	4.642	1.938	5.175	

Source: Naghavi and Yu [8].

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# CHAPTER 6 RESULTS OF MODEL APPLICATION

In this chapter, flood peaks are estimated using the eight ungauged watershed models for the Ward Creek watershed. The model parameters are estimated based on information of watershed characteristics, land use, and regional climatology. No parameter is calibrated using stream gauging data.

### 1. The Neely Model

The Neely model consists of four model parameters: (1) drainage area, A, in square miles, (2) average annual precipitation, P, in inches, (3) main channel slope, S, in ft/mile, and (4) urban adjustment ratio,  $R_L$ . These parameters have been described in detail in Chapter 5. For the Ward Creek watershed, A=4.63 square miles, P=56 inches/year, and S=3.564 ft/mile [3]. The value of  $R_L$  is estimated based on the assumption that 95 percent of the drainage area is served by storm sewer systems and the average residential lot size is 1/3 acre. The percentage of impervious area is calculated using the method described in the TR-55 report [6],

$$A_i$$
 = 0.30(Resi. area) + 0.85 (Commerc. area) + 0.72 (Industr. area)  
= 0.30 (1692) + 0.85 (919) + 0.72 (64)  
= 1335 acres (45% of the total) (6-1)

Using Figure 15 in Neely [2], the urban adjustment ratio  $(R_L)$  for the Ward Creek watershed is determined to be 3.6. The flood peaks for a corresponding rural watershed calculated from Equations (5-1a) through (5-1f) are 299.6, 442.5, 548.9, 680.6, 762.3, and 863.6 cfs, respectively. The 2-, 5-, 10-, 25-, 50-, and 100-year flood peaks for the Ward Creek watershed determined from Equations (5-2a) through (5-2f) are 1079, 1544, 1846, 2240, 2520, and 2812 cfs, respectively.

### 2. The Lowe Model

The Lowe model [3] has the same parameters as the Neely model, but they have different regression coefficients. The flood peaks for a corresponding rural watershed calculated using the Lowe model defined by Equations (5-3a) through

(5-3f) are 284.0, 444.0, 575.5, 794.4, 884.1, and 972.9 cfs, respectively. With the urban adjustment ratio  $R_L = 3.6$ , the flood peaks at the six frequencies for the Ward Creek watershed calculated using Equations (5-2a) through (5-2f) are 1022, 1478, 1781, 2208, 2481, and 2753 cfs, respectively.

### 3. The USGS Seven-Parameter Model

The seven parameters in this model are: (1) drainage area, A, in square miles, (2) basin development factor, BDF, (3) flood peak,  $Q_x$ , for a corresponding rural watershed in cfs, (4) main channel slope, S, in ft/mile, (5) the two-hour, two-year rainfall,  $I_2$ , in inches, (6) basin storage as determined by the percentage of drainage area occupied by water, ST, and (7) the percentage of drainage area occupied by impervious area,  $A_i$ .

For the Ward Creek watershed, the drainage area is 4.63 square miles (12.0 km²), and the basin development factor is 12. The flood peaks for a corresponding rural watershed estimated by the Lowe model are  $\rm Q_2=284.0$ ,  $\rm Q_5=444.0$ ,  $\rm Q_{10}=575.5$ ,  $\rm Q_{25}=794.4$ ,  $\rm Q_{50}=884.1$ , and  $\rm Q_{100}=972.9$  cfs. The main channel slope is 3.564 ft/mile. The two-hour, two-year rainfall estimated using the I-D-F relations [11] is 2.682 inches. The watershed has no permanent water body such as lakes or reservoirs. Therefore, the basin storage, ST, is 0. The impervious area is 45 percent of the watershed drainage area as shown in Equation (6-1). The flood peaks at the return periods of 2, 5, 10, 25, 50, and 100 years for the Ward Creek watershed calculated using the seven-parameter USGS model are 1233, 1716, 2122, 2631, 2901, and 3264 cfs, respectively.

#### 4. The USGS Three-Parameter Model

The parameters in the USGS three-parameter model are defined the same as in the seven-parameter model. The 2-, 5-, 10-, 25-, 50-, and 100-year flood peaks for the Ward Creek watershed calculated using Equations (5-5a) through (5-5f) are 1125, 1597, 1841, 2282, 2465, and 2732 cfs, respectively.

### 5. The Lee Model

The Lee model and the Neely model have the same parameters but different regression coefficients. The flood peaks for a corresponding rural watershed

calculated using the Lee model defined by Equations (5-6a) through (5-6f) are 450.2, 645.5, 777.1, 936.9, 1061.3, and 1172.2 cfs. The flood peaks at the return periods of 2, 5, 10, 25, 50, and 100 years for the Ward Creek watershed determined using the Neely model defined by Equations (5-2a) through (5-2f) are 1621, 2310, 2747, 3318, 3740, and 4154 cfs, respectively.

### 6. The SCS Model

To use the SCS model for estimating flood peaks from the Ward Creek watershed, the following eight parameters are needed: (1) Drainage area, A, is 4.63 square miles; (2) Rainfall peaks in Baton Rouge for 24-hour duration and at the return periods of 2, 5, 10, 25, 50, and 100 years are determined to be 4.66, 6.54, 7.61, 9.52, 11.27 and 13.05 inches, respectively, using the LADOTD 24-hour rainfall frequency maps [12]; (3) Percentage of pond and swamp areas, Aw, is zero; (4) Pond and swamp adjustment factor, Fp, is 1.0 for  $A_w = 0$ ; (5) The composite curve number, CN. As shown in Figure 6.1, the Ward Creek watershed has three types of soil, Type C, Type D and combination of Types C and D (designated as C/D in Figure 6.1) based on the SCS soil classification [6]. The composite CN for the Ward Creek watershed shown in Table 6.1 is 86; (6) Rainfall time distribution type. From Figure B-2 in the TR-55 manual it was found that the rainfall distribution type in Louisiana is III; (7) Initial abstraction, Ia, is found to be 0.325 for CN = 86 by using Table 4.1 in the TR-55 manual; and (8) Time of concentration is estimated as  $T_c = 3.0$  hours (see Appendix 2). The flood peaks using the SCS method are listed in Table 6.2. In Table 6.2, P is the 24-hour rainfall for different return periods, R is mass direct runoff calculated from equation (5-7), and  $Q_p$  is the peak discharge at different return periods determined from Equation (5-8) with  $q_u = 154$  cfs/mile<sup>2</sup>/inch.

## 7. The Louisiana Regional GEV-PWM Model

Since Ward Creek is located in the southeast region of Louisiana, the mean annual maximum discharge can be estimated by using Equation (5-9a). With A=4.63 square miles, the mean annual maximum discharge calculated using Equation (5-9a) is 1022 cfs. By multiplying the mean annual maximum discharge and dimensionless flood quantiles listed in Table 5.1, the flood peaks for the

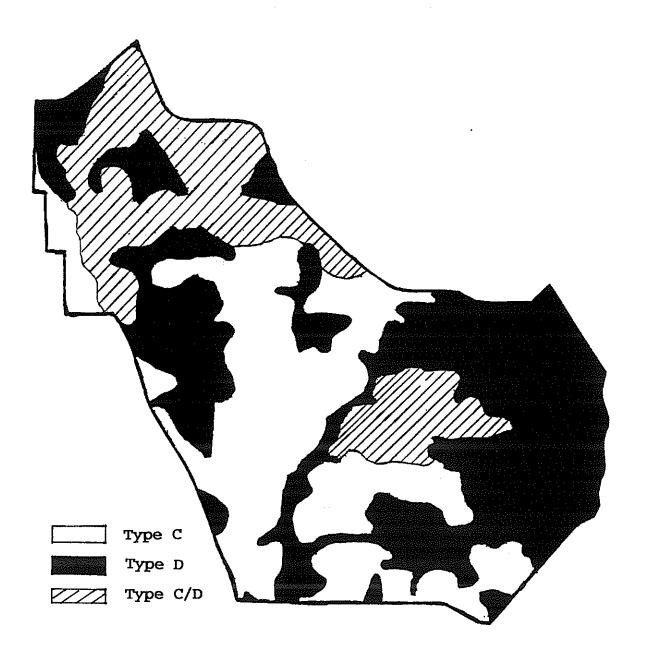


Figure 6.1 Soil classification for the Ward Creek watershed

Ward Creek watershed at the return periods of 2, 5, 10, 25, 50, and 100 years are 829, 1441, 1962, 2696, 3308, and 4002 cfs respectively.

Table 6.1
Computation of the runoff curve number for the Ward Creek watershed

Land use	Soil type	CN	Area (acres)	CN × Area
Residential *	C	82	854	69255
	C/D	84	318.3	26578
	D	87	519.8	44703
Commercial	C	94	464.1	43625
	C/D	94	172.8	16243
	D	95	282.1	26800
Industrial	C	91	32.3	2939
	C/D	92	12.0	1104
	D	93	19.6	1823
Undeveloped	C	74	145.4	10760
	C/D	77	54.1	4166
	D	80	88.4	7072
Σ			2963	255067
CN = 255067	/2963 = 86			

Note:\* Assuming the average residential lot size is 1/3 acre.

Table 6.2 Flood calculation for the Ward Creek watershed using the SCS model

	Return period (years)						
	2	5	10	25	50	100	
P (in)	4.66	6.54	7.61	9.52	11.27	13.05	
R (in)	3.151	4.924	5.954	7.811	9.527	1-1.28	
Q <sub>p</sub> (cfs)	2247	3511	4245	5570	6793	8044	

### 8. The Louisiana Regional GEV-OPT Model

The mean annual maximum flood for the Louisiana Regional GEV-OPT model is 1022 cfs, the same as that for the Louisiana Regional GEV-PMW model. By using the dimensionless flood quantiles listed in Table 5.2, the flood peaks at the return periods of 2, 5, 10, 25, 50, and 100 years calculated using the Louisiana Regional GEV-OPT model are 830, 1344, 1975, 2664, 3232, and 3849 cfs, respectively.

# CHAPTER 7 MODEL EVALUATION USING OBSERVED DATA

The flood peaks for the Ward Creek watershed are determined using the eight models in Chapter 6. The Ward Creek watershed is treated as if it were an ungauged watershed in the previous chapter. In this chapter, the flood magnitudes will be determined using the long-term stream gauging records of the watershed by following the WRC procedure [1]. The results determined from the eight models will be compared with those derived using the WRC procedure.

### 1. Water Resources Council Procedure

The flood magnitude and frequency are determined by following the Water Resources Council procedure [1].

### Step 1: Data Preparation

The annual peak discharges in Ward Creek at Government Street from 1954 through 1967 were recorded by USGS. The USGS did not establish the stage-discharge relation at this gauge station for the period from 1969 through 1995. As part of this study, a flowmeter (American Sigma 950 Area Velocity Flowmeter) was installed in Ward Creek to collect both flow velocity and water level. The observed discharge in Ward Creek is calculated using the following equations,

$$A = (B1 + ZD)D \tag{7-1}$$

$$Q = VA \tag{7-2}$$

where  $A = flow area (ft^2)$ 

B1 = bottom width (ft)

Z = side slope

D = water depth (ft)

Q = water discharge (cfs)

V = flow velocity (ft/sec)

The observed water depth and discharge data at the Ward Creek gage station are listed in Appendix 3. The rating curve for the station is represented by the following regression equations,

$$Q = 8.96D^2 + 72.70D - 13.06 \tag{7-3}$$

$$Q = 8.96(H - H_0)^2 + 72.70(H - H_0) - 1$$
 (7-4)

where D = water depth (ft)

H = stage elevation (ft, NGVD) observed by USGS at the Ward Creek gauge station

 $H_0$  = channel bottom elevation (ft, NGVD), which is 31.80 ft.

The stage-discharge relation shown in Figure 7.1 is determined by 591 data points collected from May to August 1996, which are listed in Appendix 3. The annual peak discharges from 1969 through 1995 were obtained by converting the annual peak stages to discharges using the rating curve defined by Equation (7-4). The annual maximum discharges from 1954 to 1967 were observed by the USGS. To fill the data gap in 1968 due to channel lining construction, a regression relationship between stages of Ward Creek at Government Street and at Siegen Lane, about 7 miles (11.3 km) downstream, is obtained using the USGS stage data at the two gauge stations from 1954 through 1967 (Figure 7.2). The peak stage at Siegen Lane in 1968 was 14.69 ft. The peak stage at Government Street in 1968 was determined to be 11.66 ft using the relation shown in Figure 7.2. The peak discharge in Ward Creek at Government Street in 1968 is calculated as 1850 cfs using the rating curve for Ward Creek at Government from 1961 through 1966 (see Figure 7.1).

### Step 2: Data List

Annual peak dicharge data of 42 years (1954-95) are listed in Table 7.1 and their logarithms, squares and cubes of logarithms are calculated.

### Step 3: Statistical Analysis

Using the data in Table 7.1 and equations in the WRC report [1], the mean, standard deviation and skew coefficient of these data are computed.

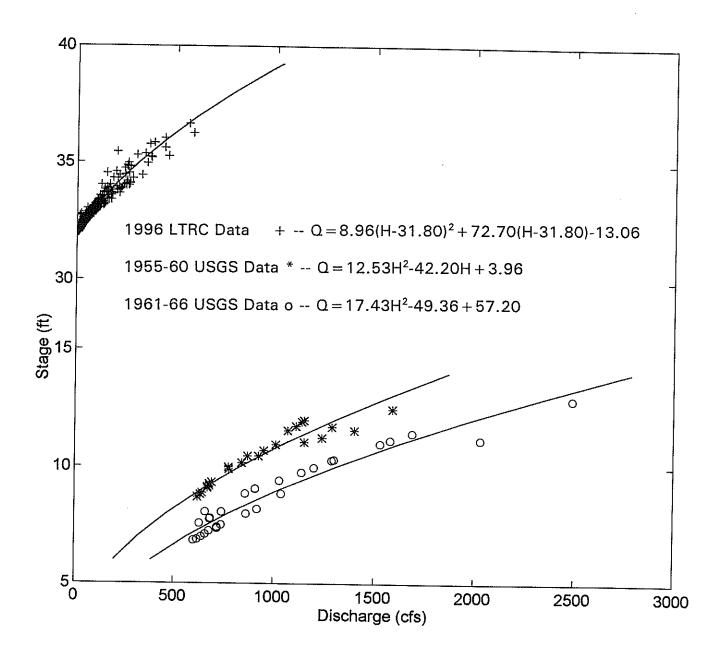


Figure 7.1
Rating curves for the Ward Creek watershed

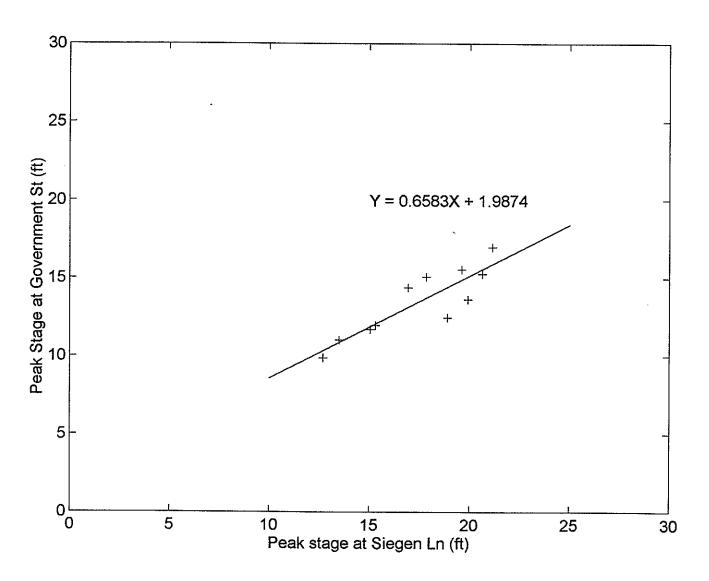


Figure 7.2
Relation between peak stages in Ward Creek at Government Street and Siegen Lane

Table 7.1 Computation of summations

Year	Peak discharge (cfs)	X=LOG(Q	) X <sup>2</sup>	X <sup>3</sup>
1954 1955 1956 1957 1958 1957 1961 1962 1963 1964 1965 1966 1967 1977 1977 1977 1978 1984 1985 1988 1988 1988 1988 1988 1988 1988	895 1150 777 1590 1290 1400 1150 1040 2030 1290 1690 1690 1690 1620 1630 1640 1690 1660 1690 1690 1690 1690 1690 169	2.95182 3.06070 2.89042 3.20140 3.11059 3.14613 3.06070 3.01703 3.30750 2.85733 3.11059 3.22789 3.18469 3.39620 3.26717 3.37107 3.13354 3.20140 3.27646 3.31597 3.20952 3.20683 3.16137 3.12385 3.13704 3.07918 3.19313 2.87506 2.75588 3.22011 3.07918 3.07918 3.07918 3.19313 2.87506 2.75588 3.22789 3.22011 3.32428 3.33846 3.33846 3.13672 3.25285 3.38382	8.71326 9.36787 8.35453 10.24894 9.67577 9.89812 9.36787 9.10249 10.93953 8.16435 9.67577 10.41925 10.14226 11.53417 10.67441 11.36410 9.81907 10.24894 10.73520 10.99566 10.30099 10.28373 9.99425 9.75845 9.84100 9.48136 10.19604 8.26598 7.59485 10.36910 9.48136 10.36910 9.48136 10.36910 11.05085 10.36910 11.05085 10.36910 11.05085 10.36910 11.14529 9.83902 11.14529 11.14529 11.14529 9.83902 11.45021	25.72000 28.67222 24.14812 32.81094 30.09735 31.14076 28.67222 27.46252 36.18245 23.32826 30.09735 33.63217 32.29997 39.17234 34.87513 38.30915 30.76840 32.81094 35.17348 36.46128 33.06117 32.97814 31.59549 30.48395 30.48395 30.87160 29.19482 32.55724 23.76519 20.93045 33.38961 29.19482 28.56576 22.87847 33.63217 33.38961 29.19482 28.56576 22.87847 33.63217 33.38961 36.73616 35.89954 37.20807
		104.04000	+13.2343U	1329.40200

$$X_m = \frac{\Sigma X}{N} = \frac{132.54690}{42} = 3.15588$$

$$S^2 = \frac{\sum X^2 - (\sum X)^2 / N}{N-1}$$

$$=\frac{419.29490 - (132.54690)^2/42}{42-1}$$

= 0.024219

S = 0.15562

$$G = \frac{N^{2}(\Sigma X^{3}) - 3N(\Sigma X)(\Sigma X^{2}) + 2(\Sigma X)^{3}}{N(N-1)(N-2)S^{3}}$$

$$42(42-1)(42-2) \times 0.15562^3$$

= -0.74223

where X = logarithm of annual peak discharge

N = number of observations of the data set

 $X_m$  = mean of log-transformed annual maximum series

S = standard deviation of logarithms

G = skew coefficient of the log-transformed flood series.

## Step 4: Computation of the Frequency Curve Coordinates

The frequency curve coordinates are calculated using the follwing equation,

$$LOG Q = X_m + KS \tag{7-5}$$

where K is the frequency factor, which is a function of the skew coefficient and the selected frequency. The K values at the six frequencies for Ward Creek

watershed for S=0.15562 obtained from Appendix 3 of the WRC report [1] are listed in Table 7.2. The flood peaks are computed using Equation (7-5). An example computation for an exceedance of 0.5 (two-year occurrence) is shown below, and others are list in Table 7.2.

Log Q = 
$$X_m$$
 + KS  
= 3.15588 + 0.12263 × 0.15562  
= 3.17496

Q = 1496 cfs

The peak discharge at a selected frequency can be determined by following four steps: (1) find the logarithm of annual peak discharge; (2) calculate  $X_m$ , S and G; (3) find the value of K from Appendix 3 of the WRC report [1] for the frequency and skew coefficient; and (4) calculate discharge using Eq. (7-5).

### 2. Evaluation of the Model Results

The flood peaks for the Ward Creek watershed calculated in Chapter 6 from the eight models are summarized in Table 7.3 and plotted in Figure 7.3 along with the observed data from the WRC procedure. The eight models are evaluated by comparing the flood peaks calculated using the models with the flood peaks derived from historical flood records using the WRC procedure. The relative root mean square error (RRMSE) is used to evaluate the flood prediction models.

$$RRMSE = \sqrt{\frac{1}{n} \sum_{Q_{WRC, X}} \left[ \frac{Q_x - Q_{WRC, X}}{Q_{WRC, X}} \right]^2}$$
 (7-6)

where RRMSE = relative root mean square error

 $Q_x$  = flood peak at the return period of x years calculated using one of the eight models

 $Q_{\text{WRC},x}$  = flood peak at the return period of x years calculated using the WRC procedure

n = number of frequencies to be compared (n = 6 here).

Table 7.2 Computation of frequency curve coordinates

Return Period (years)	Freq. (1/yr.)	К	Log(Q)	Q (cfs)
2	0.5	0.12263	3.17496	1496
5	0.2	0.85662	3.28919	1946
10	0.1	1.17598	3.33889	2182
25	0.04	1.40871	3.37510	2372
50	0.02	1.63909	3.41096	2576
100	0.01	1.77518	3.43213	2705

Table 7.3
Results of model application to the Ward Creek watershed

		Return period (years)					
Model	2	5	10	25	50	100	RRMSE
Neely	1079	1544	1846	2240	2520	2812	0.16
Lowe	1022	1478	1781	2208	2481	2753	0.18
USGS-7	1233	1716	2122	2631	2901	3264	0.14
USGS-3	1125	1597	1841	2282	2465	2732	0.14
Lee	1621	2310	2747	3318	3740	4154	0.35
scs	2247	3511	4245	5570	6793	8044	1.30
GEV-PWM	829	1441	1962	2696	3308	4002	0.32
GEV-OPT	830	1344	1975	2664	3232	3849	0.30
WRC	1496	1946	2182	2426	2576	2705	

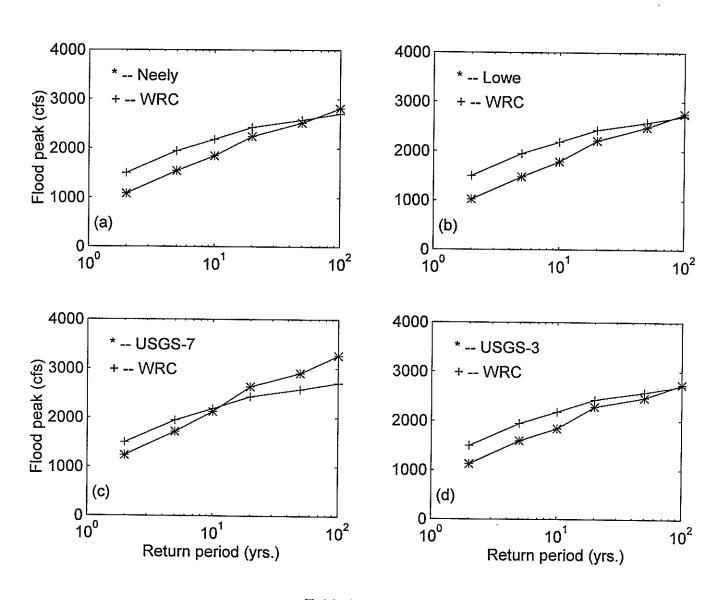


Table 7.3
Comparison of model results with those using the WRC procedure

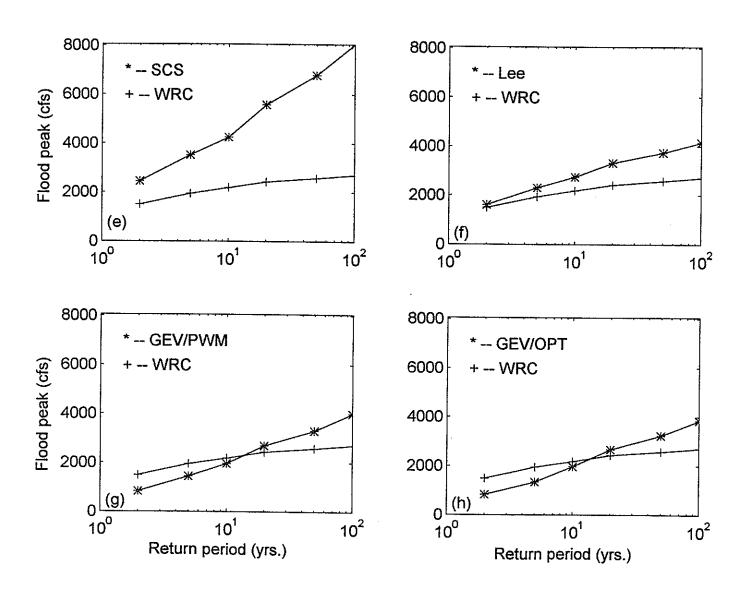


Table 7.3 (Continued)
Comparison of model results with those using the WRC procedure

The values of RRMSE for the eight models are listed in Table 7.3. The USGS 7-parameter model, USGS 3-parameter model, the Lowe model and the Neely model have the least relative root mean square error. Therefore, these four models are better than the other models in predicting flood frequency for the Ward Creek watershed. The four models have better accuracy because: (1) the field data for developing the Lowe model and the Neely model were all from Louisiana, and one watershed was used for development of the two USGS models; (2) the drainage area of the Ward Creek watershed is within the ranges which the four models apply; and (3) the parameters for the four models, such as drainage area, main channel slope, mean annual precipitation, impervious area, basin storage, basin development factor, urban adjustment ratio, etc., do not depend significantly on personal judgement.

The Louisiana regional GEV-PWM and GEV-OPT models were developed based on field data for rural watersheds. The accuracy of flood prediction for highly developed urban watersheds like the Ward Creek watershed is relatively low. The GEV-PWM and GEV-OPT models are simple to apply because only drainage area and watershed location are needed. However, they can be used for quick but rough estimation of flood peaks. For more accurate estimation of flood magnitudes, more information of the watershed is needed to use other models such as the two USGS models, the Lowe model, and the Neely model.

The SCS model has the least accuracy for the Ward Creek watershed. The accuracy of the SCS model is low at all six frequencies. This is due to the limitations of the model itself and errors in estimating the model parameters. When the runoff curve number is estimated, personal judgement is involved. When estimating the unit peak discharge, the ratio of I<sub>a</sub>/P was out of the range in the nomograph [6] at all six frequencies. The TR-55 manual [6] suggests that the boundary values be used. Therefore, the unit peak discharge (q<sub>u</sub>) was determined to be 154 for all six frequencies. Obviously, a large error was introduced.

- <del>\*\*</del>

### **CHAPTER 8**

# DETERMINATION OF FLOOD PEAKS USING SHORT-TERM FIELD DATA CALIBRATION

The SCS model is re-applied here with parameters calibrated using short-term field data. Two parameters, runoff curve number (CN) in Equation (5-7) and unit peak discharge  $(q_u)$  in Equation (5-8), in the SCS model are calibrated using field data. However, parameter calibration is not applicable to the rest of the eight models because they were developed using data of multiple watersheds.

### 1. Runoff Curve Number Calibration

A data set of 24-hour rainfall and net runoff was obtained from the rain gauge and flowmeter records. In Table 8.1, 24-hour rainfall is the arithmetic average of 24-hour rainfall recorded by the three rain gauges. To calculate the net runoff for a rainfall event, the stage data corresponding to the rainfall event are converted to discharge data using the rating curve derived in Chapter 7. The net runoff is calculated by integration. Baseflow is neglected because the depth of baseflow in Ward Creek at Government Street is usually one to two inches. The 24-hour rainfall and net runoff data listed in Table 8.1 will be used for runoff curve number calibration. The standard deviation of the calculated net runoff using Equation (5-7) from the observed net runoff is represented by the following equation

$$s = \sqrt{\frac{1}{(n-1)} \sum_{i=1}^{n} (R_{ci} - R_{oi})^2}$$
 (8-1)

where s = standard deviation

n = number of storm events

 $R_{\text{ci}}\,=\,\text{calculated}$  net runoff (in.) during the i'th storm event

 $R_{\text{oi}}\,=\,\text{observed}$  net runoff (in.) during the i'th storm event.

 $R_{\text{ci}}\,\text{is}$  calculated using Equation (5-7). By substituting Equation (5-7) into equation (8-1), we obtain

$$S = \sqrt{\frac{1}{n-1} \sum_{i=1}^{n} \left( \frac{[P_i - 0.2 \ S)]^2}{P_i + 0.8 \ S} - R_{oi} \right)^2}$$
 (8-2)

where S = 1000/CN - 10.

Table 8.1

Rainfall and runoff field data for the Ward Creek watershed

		Rainfal	l (inches)		Net runoff	Peak
Date	Foster	Mission	Scenic	Average	(inches)	runoff (cfs)
03/09/94	0.94	1.91	3.11	1.99	0.47	324
06/12/94	2.79			2.79	0.79	985
07/08/94	3.16	3.19		3.18	0.66	1095
07/27/94	1.90	2.39	2.18	2.16	0.81	1383
08/23/94		1.42	1.79	1.61	0.40	870
03/07/95	1.77	1.82	1.82	1.8	0.53	768
03/13/95	3.57	3.43	3.54	3.51	1.35	898
03/28/95	3.50	3.89	4.20	3.86	1.52	1024
04/10/95	6.10	6.60	6.43	6.38	4.12	2449
05/18/95	2.90	3.10	4.17	3.39	1.33	1353
09/21/95	1.38	1.84	2.09	1.77	0.43	524
11/02/95	4.80	4.37		4.59	0.85	827
01/01/96	0.74	1.00	1.19	0.98	0.24	342
01/26/96	2.77	3.55	3.77	3.36	1.37	1525
02/28/96	1.76	2.02	1.79	1.86	0.16	333
04/12/96	2.22	1.85	2.66	2.24	0.78	1007
06/24/96	1.35	1.73	1.26	1.45	0.54	1019

CN is determined by minimizing the standard deviation in Equation (8-2). This is done by: (1) plugging the field-observed values of  $P_i$  and  $R_{oi}$  in Equation (8-2); (2) finding the derivative of s with respect to CN; (3) setting ds/d(CN) = 0; and (4) solving the equation to determine CN. An alternative method to determine CN is by: (1) calculating the values of standard deviation using Equation (8-2) for CN's from 40 to 100 with an interval of 5; (2) plotting standard deviation (s) vs. runoff curve number (CN); and (3) finding the value of CN at which the standard deviation is minimum. Figure 8.1 shows the relation between standard deviation and runoff curve number. The value of standard deviation becomes minimum when deviation deviation

### 2. Unit Peak Discharge

The unit peak discharge in Equation (5-8) is calibrated using the rainfall-runoff data listed in Table 8.1. The unit peak discharge is defined as

$$Q_{u} = \frac{Q_{p}}{ARF_{p}} \tag{8-3}$$

where  $q_u$ , A, R,  $F_p$  and  $Q_p$  are defined in Chapter 5. Using the field data listed in Table 8.1, a regression relation between  $q_u$  and  $(ARF_p)$  is developed as follows

$$q_{u} = \frac{632}{(A R F_{p})^{0.25}} - 192 \tag{8-4}$$

The curve fitting is shown in Figure 8.3. The unit peak discharge decreases with A R  $F_{\rm p}$ .

## 3. Model Application Using Calibrated Parameters

Table 8.2 lists the flood peaks calculated from the SCS model using the calibrated runoff curve number and unit peak discharge, in which P is the 24-hour rainfall at a certain return period derived from the LADOTD 24--hour rainfall frequency maps [12]; R is the net 24-hour runoff at that return period calculated

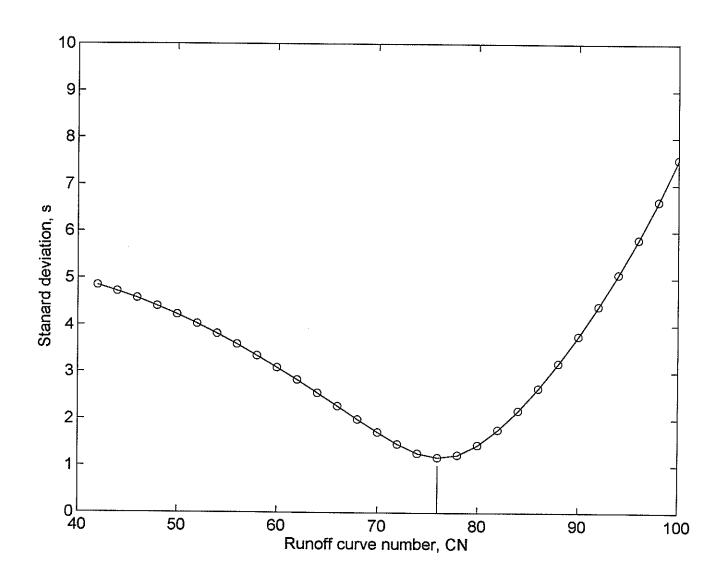


Figure 8.1 Calibration of runoff curve number for the Ward Creek watershed

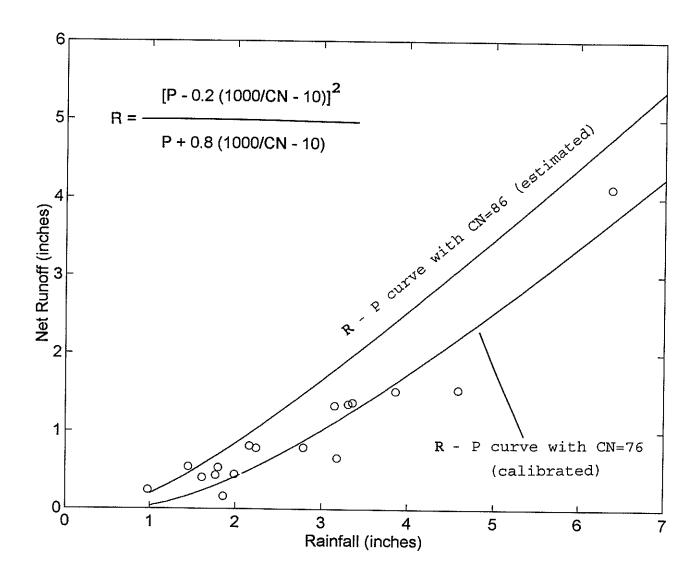


Figure 8.2
Comparison of the estimated runoff curve number with the calibrated runoff curve number

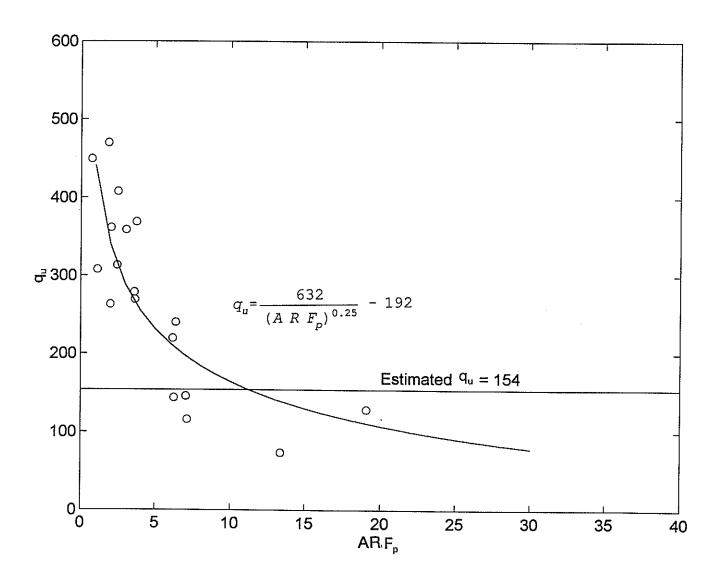


Figure 8.3
Calibration of unit peak discharge for the Ward Creek watershed

using Equation (5-7) with the curve number equal to 76; and  $q_u$  is the unit peak discharge calculated using Equation (8-4). The peak runoff rates,  $Q_p$ , at the return periods of 2, 5, 10, 25, 50, and 100 years calculated using Equation (5-8) are 1667, 2060, 2203, 2345, 2377, and 2333 cfs, respectively.

Table 8.2
Flood calculation using the SCS model with parameter calibration

	Return period (years)						
	2	5	10	25	50	100	
P (in)	4.66	6.54	7.61	9.52	11.27	13.05	
R (in)	2.258	3.851	4.804	6.558	8.203	9.901	
$q_u$	190	135	110	86.1	72.4	59.7	
Q <sub>ρ</sub> (cfs)	1667	2060	2203	2345	2377	2333	

Compared with the flood peaks derived using the WRC procedure, the relative errors of the flood peaks determined using the SCS model with parameter calibration are 11.4, 5.9, 1.0, -3.3, -7.7, and -13.8 percent, respectively. The relative errors of the flood peaks calculated using the SCS model without parameter calibration (in Table 7.3) are 50.2, 80.4, 94.5, 129.6, 163.7, and 197.4 percent, respectively. The RRMSE defined in Equation (7-6) for the SCS method with parameter calibration is 0.070, compared with 1.30 for the SCS model without parameter calibration. Overall, flood prediction by the SCS model using short-term field data calibration can be significantly improved as compared with the parameter estimation procedure suggested by the TR-55 manual [6]. The problem is that the calculated 50-year flood is higher than 100year flood. This is because of the error in parameter calibration at 50- and 100year return periods. Since the most severe rainfall storm used for parameter calibration in this study is equivalent to five-year return period (see Table 8.1), the calculated flood peaks at 2-, 5-, 10-, and 25-year are more reliable. In case the 50-year or 100-year flood is needed for a design application, the 50-year and 100-year floods may be corrected by extrapolation. In Figure 8.4, the regression relation between flood peak and return period is determined using

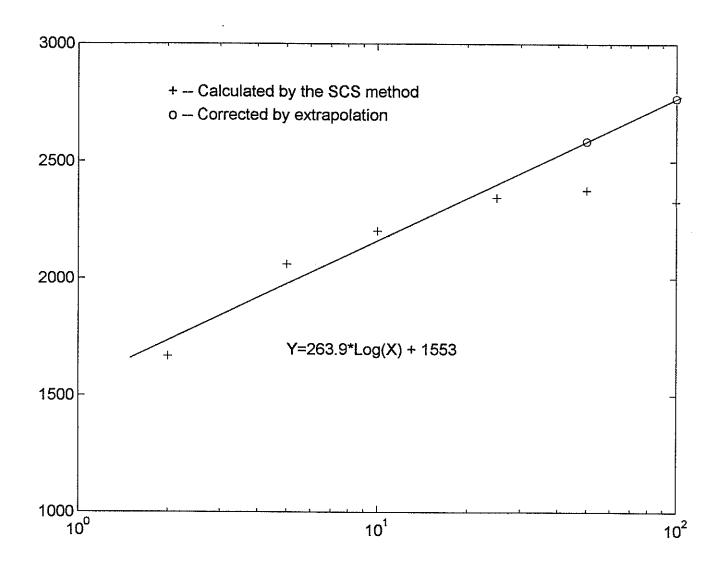


Figure 8.4 Correction of 50- and 100-year flood peaks by extrapolation

four points at 2-, 5-, 10-, and 25-year return periods. The corrected flood peaks at 50- and 100-year return periods determined using extrapolation are 2585 and 2768 cfs, respectively, which are closer to the results determined using the WRC procedure (see Table 7.3).

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# CHAPTER 9 CONCLUSION

In this study, eight flood prediction models for ungauged watersheds are applied to the Ward Creek watershed, and the results are compared with those derived using long-term stream gauging records by following the U.S. Water Resources Council procedure. It is found that the USGS 7-parameter model, the USGS 3-parameter model, the Lowe model, and the Neely model have better accuracy than the others. The Lee model was developed using watershed data from Louisiana, Mississippi and Arkansas, however, the model overestimates flood magnitudes for the Ward Creek watershed. The Louisiana Regional GEV-PWM and the Louisiana regional GEV-OPT models are simple to apply because only drainage area and watershed location are needed. They may be used for quick and rough estimation of flood peaks. The SCS model is most widely used for flood prediction. However, the accuracy of flood prediction for the Ward Creek watershed using the SCS model is low. The RRMSE defined in Equation (7-6) is as high as 1.30 and the relative errors of the peak discharge prediction at all six frequencies are 50.2, 80.4, 94.5, 129.6, 163.7, and 197.4 percent, respectively. Errors in parameter estimation are substantial because the model was not calibrated using local data.

This study demonstrates that the accuracy of peak discharge prediction using the SCS model can be significantly improved if parameters are calibrated using short-term field data. With the runoff curve number and unit peak discharge calibrated using two-year rainfall-runoff data for the Ward Creek watershed, the RRMSE defined in Equation (7-6) is reduced to 0.07 and the relative errors of the peak discharge prediction at all six frequencies are 11.4, 5.9, 1.0, -3.3, -7.7, and -13.8 percent, respectively.

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# CHAPTER 10 RECOMMENDATION

To estimate flood peaks from a watershed, it is of vital importance to choose an appropriate model of flood prediction. The selection of a flood prediction model depends on the watershed data availability and required accuracy.

- 1. For quick and approximate estimation of flood peaks with known drainage area and watershed location in Louisiana, the Louisiana Regional GEV-PWM or the GEV-OPT model may be used.
- 2. If the watershed geometry, land coverage and local climatic data are known, one of the four models (the two USGS models, the Lowe model and the Neely model) may be used. One may also use all the four models and take the average of the model results.
- 3. If accurate flood prediction is technically necessary and economically justified, a temporary network of rain gauges and flowmeter may be set up to obtain rainfall-runoff data for one to two years. The SCS model is be applied with the parameters calibrated using the short-term field data. This study has shown that the accuracy of the SCS model can be significantly improved with the runoff curve number and unit peak discharge calibrated using short-term rainfall-runoff data. The procedure is (1) calibrate runoff curve number using 24-hour rainfall and net runoff data; (2) develop a unit peak discharge curve using peak discharge and net runoff data; (3) find 24-hour rainfall at the return periods of 2, 5, 10, 25, 50, and 100 years using 24-hour rainfall frequency maps or I-D-F curves; (4) calculate net runoff at each frequency using Equation (5-7) in this report with calibrated runoff curve number; (5) determine unit peak discharge at each frequency using the calibrated unit peak discharge curve; and (6) compute peak discharge at each frequency using Equation (5-8) in this report.
- 4. If a watershed has long-term stream gauging data but has undergone significant hydrological changes, the flood peaks may be determined using the SCS

model with parameter calibrated using the most recent stream gauging data and local rainfall data.

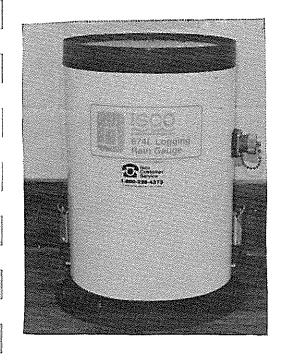
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## **APPENDIX 1**

## PICTURES OF RAIN GAUGE AND FLOWMETER



(a) ISCO 674L rain gauge

(b) American Sigma 950 flowmeter



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## **APPENDIX 2**

## TIME OF CONCENTRATION CALCULATION FOR THE WARD CREEK WATERSHED

Time of concentration can be estimated by the TR-55 method [6]. Figure A.2 shows the ranges of three types of surface flows (sheet flow, shallow concentrated flow, and open channel flow) and locations of eight cross-sections along the main channel of Ward Creek.

For sheet flows, the following equation [13] is used,

$$T_c = \frac{0.42 (nL)^{0.8}}{P_2^{0.5} S^{0.4}} \tag{A-1}$$

where the Manning's roughness coefficient (n) was chosen as 0.05 considering that the drainage area is composed of paved surface and short-grass surface, the overland flow length (L) is 300 feet, the average slope (S) is 0.001, and the 2-year 24-hour rainfall ( $P_2$ ) is 4.66 inches determined using Louisiana 24-hour rainfall frequency maps [12]. The travel time of the sheet flow is 26.8 minutes as calculated using Equation (A-1).

For shallow concentrated flows, two equations were developed based on the Chezy-Manning equation with different assumptions of the Manning roughness coefficient (n) and hydraulic radius ( $r_h$ ). With the assumptions of the Manning roughness coefficient n=0.05 and hydraulic radius  $r_h=0.4$  ft for unpaved surfaces, and n=0.025 and  $r_h=0.2$  ft for paved surfaces, the following equations are used to estimate the time of concentration,

$$T_c = \frac{L/60}{16.1345\sqrt{s}} \quad unpaved \tag{A-2}$$

$$T_c = \frac{L/60}{20.3282\sqrt{s}} \quad paved \tag{A-3}$$

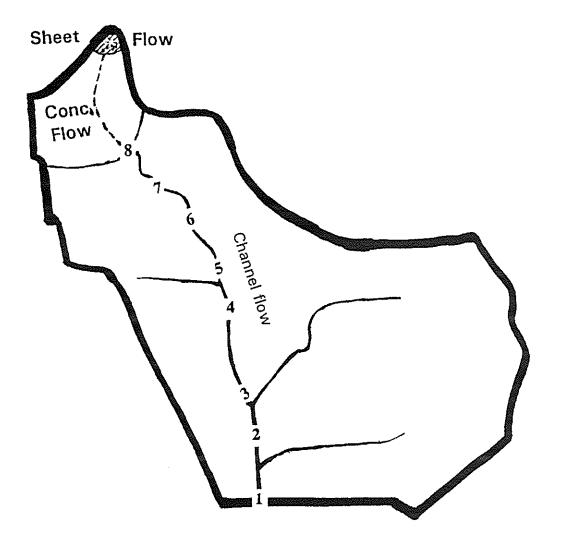


Figure A.2
Flow path for calculating time of concentration for the Ward Creek watershed

As shown in Figure A.2, the length of the shallow concentrated flow for the Ward Creek watershed is 4000 ft and the average land slope is 0.00127. By taking the average values of the time of concentration computed from Equations (A-2) and (A-3), the travel time for the shallow concentrated flow is 104.0 minutes.

To estimate travel time for channel flow, the main channel is divided into eight sections. Figure A.2 shows the locations of the eight cross-sections and Table A.1 lists the geometric parameters of each cross-section. The discharge at cross-section 8 is estimated by using the Rational Formula. The drainage area above cross-section 8 is about 256 acres. The runoff coefficient is estimated as 0.35 based on single-family residential area. The two-year rainfall itensity during time of concentration of 26.8 + 104.0 = 131 minutes computed using the I-D-F curve in the LADOTD Hydraulcs Manual [11] is I = 1.355 in/hr. The discharge at cross-section 8 computed using the Ration Formula is 121.3 cfs. The mean annual maximum discharge at cross-section 1 is estimated using Equation (5-9a) at 1022 cfs. Water discharges at other cross-sections were calculated using the total channel length (including tributaries) above the cross-section as the weight by the following equation:

$$Q_{i} = Q_{8} + \frac{\sum_{j=i}^{8} L_{j}}{\sum_{j=1}^{8} L_{j}} (Q_{1} - Q_{8})$$
(A-4)

where  $L_j$  is the total channel length including tributaries between two successive cross-sections, j and j+1.

Flow velocity between two cross-sections is calculated using the Chezy-Manning equation by using average channel geometries and flow discharges. As a result, the travel time through the main channel is 50.9 minutes. The result of the calculation is listed in Table A.1. The time of concentration of the entire watershed is 26.8 + 104.2 + 50.9 = 181.9 minutes or approximately 3.0 hours.

Table A.1
Estimation of travel time through the main channel of Ward Creek

Cross-section								
no.	1	2	3	4	5	6	7	8
Station no.	7 + 12.5	32+ 1.0	41 + 95.2	73+ 50	87+ 17.5	110+ 50	124+ 78	151 +79
Discharge (cfs)	1022	821	507	419	302	237	197	121
b (ft)	24	18	18	14	12	8	8	8
Z (ft/ft)	2	1.5	1.5	1.5	1.5	1.5	1.5	1.5
E1 (ft)	31.02	33	33.95	36.63	37.67	41.08	42.22	50
E2 (ft)	44	46	49	48.5	49	52	53	54.5
Distance (ft)	<b></b>	2488	995	3155	1388	2333	1428	2701
Tributary length (ft)	4724	10236	0	2835	0	0	0	0
Slope (%)		0.08	0.095	0.085	0.076	0.146	0.08	0.288
Avg. b		21	18	16	13	10	8	8
Avg. n	-	0.016	0.016	0.016	0.016	0.016	0.016	0.025
Avg. Q		921.2	663.8	463	360.4	269.2	216.7	159
Flow depth (ft)		5.11	4.37	3.91	3.87	3.11	3.57	3.93
Flow area (sq ft)	<del></del>	146.5	107.3	85.46	72.8	45.68	47.59	54.6
Velocity (ft/sec)		6.29	6.19	5.42	4.95	5.89	4.55	2.91
Travel time (minutes)		6.59	2.68	9.7	4.6	6.6	5.23	15.47

Note: Watershed information is from East Baton Rouge Parish Department of Public Works.

APPENDIX 3
STAGE AND DISCHARGE DATA FOR WARD CREEK AT GOVERNMENT ST (1996)

Month	Date	Time (CDT)	Depth (ft)	Velocity (ft/sec)		Discharge (cfs)
5 5	23 23	6:30 6:45	0.167	0.270	4.068	1.10
5 5	23 23	7:00 7:15	0.273	1.122	6.698	7.51
5	23	7:30	0.280	0.997 0.926	6.873 6.698	6.85 6.20
5 5 5 5 5	23 23	7:45 8:00	0.271 0.270	0.764 0.517	6.648 6.623	5.08 3.42
5	23	8:15	0.269	0.444	6.598	2.93
5	23	8:30	0.267	0.364	6.548	2.38
5	23	8:45	0.258	0.327	6.323	2.07
5	23	9:00	0.249	0.130	6.099	0.79
5 5 5	23 23	9:15 9:30	0.230	0.009	5.626	0.05
5	23	9:45	0.176	0.068	5.031 4.290	0.34 0.30
5	23	15:00	0.173	0.219	4.216	0.92
5	23	15:15	0.173	0.389	4.216	1.64
5	23	15:30	0.183	0.232	4.462	1.04
5	23	15:45	0.189	0.142	4.610	0.65
5	23	16:00	0.191	0.155	4.660	0.72
5	23	16:15	0.198	0.054	4.833	0.26
5	23	16:30	0.202	0.023	4.932	0.11
5	23	16:45	0.204	0.040	4.981	
5 5	23	17:00	0.200	0.108	4.882	0.53
5	23	17:15	0.206	0.064	5.031	0.32
	23	17:30	0.207	0.064	5.056	0.32
5	23	17:45	0.207	0.144	5.056	0.73
5	23	18:00	0.209	0.117	5.105	0.60
5	23	18:15	0.200	0.331	4.882	1.62
5	23	18:30	0.186	0.545	4.536	2.47
5	23	18:45	0.188	0.493	4.586	2.26
5	23	19:00	0.195	0.300	4.759	1.43
5	23 23	19:15 19:30	0.191	0.352 0.316	4.660	1.64
5 5 5 5	23	19:45	0.183	0.532	4.660 4.462	1.47
5	23	20:00	0.182	0.336	4.438	1.49
	23	20:15	0.170	0.501	4.142	2.07
5	29	12:15	0.735	1.533	18.596	28.51
5	29	12:30	0.985	1.891	25.329	47.90
5 5 5	29 29	12:45 13:00	1.331 1.769	2.301 2.412	34.991 47.792	80.51 115.27
5	29	13:15	2.025	2.498	55.569	138.81
5	29	13:30	2.108	2.593	58.137	150.75
5 5	29 29	13:45 14:00	2.067	2.779 2.842	56.866	158.03
5	29	14:15	2.000 1.906	2.646	54.800 51.927	155.74 137.40
5	29	14:30	1.769	2.676	47.792	127.89
5	29	14:45	1.609	2.716	43.042	116.90
5	29	15:00	1.441	2.556	38.146	97.50
5	29	15:15	1.300	2.485	34.109	84.76
5 5 5	29 29	15:30 15:45	1.171 1.058	2.408 1.593	30.474 27.335	73.38 43.54
5	29	16:00	0.958	2.364	24.592	58.14
5	29	16:15	0.868		22.152	51.48
5 5	29 29	16:30	0.796 0.733	2.273	20.219	45.96
5 5	29	16:45 17:00	0.679	2.229	18.543	41.33 38.15
5	29	17:15	0.628	2.191	15.777	34.57
	29	17:30	0.587	2.128	14.707	31.30
5	29	17:45	0.552	1.990	13.798	27.46
5	29	18:00	0.520	1.987	12.970	25.77
5 5 5 5 5 5 5 5 5	29	18:15	0.485	1.827	12.069	22.05
	29	18:30	0.452	1.770	11.223	19.87
5 5	29 29	18:45 19:00	0.430	1.289 1.448	10.661	13.74
5	29	19:15	0.383	0.979	9.466	14.41 9.27
5	29 29	19:30 19:45	0.367	0.312 0.212	9.061 8.606	2.83 1.82
5	29	20:00	0.324	0.308	7.976	2.46
5	29	20:15	0.308	0.103	7.574	0.78

Month	Date	Time (CDT)	Depth (ft)	Velocit (ft/sec)		Discharge
5 5	29 29	20:30	0.281	0.700	(ft2) 6.898 6.698	(cfs) 4.83 0.52
5	29	21:00	0.253	0.207	6.198	1.28
5	29 29	21:15 22:00	$0.234 \\ 0.194$	0.241 0.157	5.726 4.734	1.38 0.74
5	29 7	22:15 16:00	0.172 0.238	0.207 1.353	4.191 5.825	0.87 7.88
6 6	7 7	16:15 16:30	0.674 1.930	1.373 2.747	16.984 52.658	23.32 144.65
6 6	7 7	16:45 17:00	3.061 3.009	2.810 3.063	89.263 87.486	250.83 267.97
6 6	7 7	17:15 17:30	2.635 2.236	2.833	74.977 62.142	212.41 164.43
6 6	7 7	17:45 18:00	1.871	2.631 2.577	50.865 41.868	133.83
6	7 7	18:15 18:30	1.317	2.374	34.593	107.89 82.12
6	, 7 7	18:45	0.948	2.321	28.830 24.320	66.91 53.70
6	7	19:00 19:15	0.814	2.072	20.701 17.854	42.89 36.90
6	7 7	19:30 19:45	0.617 0.543	1.949	15.489 13.565	30.19 25.91
6	7	20:00 20:15	0.481 0.429	1.833 1.786	11.967 10.636	21.93 19.00
6 6	7 7	20:30 20:45	0.383 0.347	1.713 1.682	9.466 8.556	16.22 14.39
6 6	7 7	21:00 21:15	0.311 0.281	1.614 1.332	7.649 6.898	12.35 9.19
6 6	7 7	21:30 21:45	0.259 0.250	1.069 0.385	6.348 6.124	6.79 2.36
6 6	7 7	22:00 22:15	0.210 0.205	1.571 0.851	5.130 5.006	8.06 4.26
6 6	7 7	22:30 22:45	0.183 0.283	1.474 1.695	4.462 6.948	6.58 11.78
6	7	23:00 23:15	0.808	2.143 2.452	20.540 32.327	44.02 79.27
6	7	23:30 23:45	1.436	2.519	38.002 40.728	95.73 107.89
6	8 8	0:00 0:15	1.516	2.600	40.320	104.83
6	8	0:30	1.319	2.534	38.002 34.650	96.30 85.52
6	8	0:45 1:00	1.190	2.508	31.006 27.418	77.76 63.99
6	8	1:15 1:30	0.945	2.238 2.241	24.238 21.399	54.24 47.95
6 6	8 8	1:45 2:00	0.747 0.665	2.103 2.072	18.914 16.747	39.78 34.70
6 6	8 8	2:15 2:30	0.597 0.537	2.028 1.947	14.967 13.410	30.35 26.11
6 6	8 8	2:45 3:00	0.484 0.436	1.825 1.812	12.044 10.814	21.98 19.60
6 6	8 8	3:15 3:30	0.395 0.359	1.711 1.661	9.771 8.859	16.72 14.71
6 6	8 8	3:45 4:00	0.328 0.306	1.498 0.825	8.077 7.524	12.10 6.21
6 6	8 8	4:15 4:30	0.289 0.264	0.407 0.257	7.098 6.473	2.89
6	8	4:45 5:00	0.247	0.196	6.049 5.527	1.19
6	8	5:15 5:30	0.208 0.195	0.146	5.080 4.759	0.74 0.36
6 6	14 14	10:45 11:00	0.172	1.242	4.191	5.21
6	14	11:15	0.898	1.459	12.790 22.962	18.66 40.37
6 6	14 14	11:30 11:45	0.813	1.732	20.674	35.81 37.36
6 6	14 14	12:00 12:15	1.227	1.588	32.045 62.142	50.89 122.98
6 7 7	14 3	12:30 16:00	3.177 0.187	2.759 0.953	93.257 4.561	257.30 4.35
7	3	16:15 16:30	0.969 2.797	0.794 2.457	24.892 80.338	19.76 197.39
7 7	3 3 3 3	16:45 17:00	3.972- 4.040	3.108	121.835 124.377	364.41 386.56
7 7	3	17:15 17:30	3.567 3.007	3.202 2.914	107.014 87.418	342.66 254.74

Month	Date			Veloci		Discharge
7 7	3	(CDT) 17:45	(ft) 2.513	(ft/sec 2.888	70.996	(cfs) 205.04
7	3	18:00 18:15	2.100 1.768	2.731 2.592	57.889 47.762	158.09 123.80
7 7	3	18:30 18:45	1.499 1.288	2.495 2.331	39.826 33.769	99.37 78.72
7 7	3	19:00 19:15	1.119 0.987	2.216 2.116	29.024 25.384	64.32 53.71
7	3	19:30 19:45	0.878 0.792	2.009	22.422	45.05
7	3	20:00	0.719	1.998	20.113 18.172	40.19 34.31
7	3	20:15 20:30	0.659 0.606	1.869 1.818	16.590 15.202	31.01 27.64
7 7	3 3	20:45 21:00	0.562 0.522	1.815 1.762	14.057 13.022	25.51 22.94
7 7	3 3	21:15 21:30	0.488 0.454	1.724 1.584	12.146 11.274	20.94 17.86
7 7	3 3	21:45 22:00	0.430 0.410	1.432 0.802	10.661 10.152	15.27 8.14
7 7	33333333333333333333333	22:15 22:30	0.397 0.374	0.314 0.195	9.821 9.238	3.08 1.80
7 7	3	22:45 23:00	0.357 0.342	0.202	8.808 8.430	1.78
, 7 7	3	23:15	0.316	0.625	7.775	0.11 4.86
7	3	23:30 23:45	0.307 0.297	0.357 0.171	7.549 7.298	2.69 1.25
7 7	7	16:15 16:30	0.214 0.469	1.350 1.907	5.229 11.659	7.06 22.23
7 7	7 7	16:45 17:00	1.309 1.611	2.485 2.511	34.365 43.101	85.40 108.23
7 7	7 7	17:15 17:30	1.560 1.414	2.430 2.380	41.605 37.368	101.10 88.94
7 7	7	17:45 18:00	1.251	2.247	32.722 28.413	73.53 58.11
7 7	7	18:15 18:30	0.967	1.984	24.838	49.28
7 7	7	18:45	0.857	1.906	21.856 19.419	41.66 36.45
7	7 7 -	19:00 19:15	0.691 0.626	1.796 1.727	17.432 15.725	31.31 27.16
7 7	7 7	19:30 19:45	0.572 0.528	1.725 1.689	14.317 13.177	24.70 22.26
7 7	7 7	20:00 20:15	0.488 0.452	1.648 1.615	12.146 11.223	20.02 18.13
7 7	7 7	20:30 20:45	0.422 0.398	1.617 1.239	10.457 9.847	16.91 12.20
7 7	7 7	21:00 21:15	0.373 0.352	1.212	9.213 8.682	11.17 8.66
7 7	, 7 7	21:30	0.342	0.489	8.430	4.12
7	7	21:45 22:00	0.321 0.312	0.606 0.080	7.901 7.675	4.79 0.61
7 7	7 7	22:15 22:30	0.298 0.285	0.117 0.272	7.323 6.998	0.86 1.90
7 7	7 7	22:45 23:00	0.265 0.255	0.455 0.510	6.498 6.248	2.96 3.19
7 7	7 7	23:15 23:30	0.255 0.248	0.155 0.052	6.248 6.074	0.97 0.32
7 7	7 13	23:45 12:00	0.240 0.255	0.161	5.875 6.248	0.95
7 7	13	12:15	0.526	0.456	13.125	4.13 5.99
7	13 13	12:30 12:45	0.484 0.485	1.288	12.044 12.069	15.51 17.73
7 7 7	13 13	13:00 13:15	0.433 0.408	1.481 1.514	10.738 10.101	15.90 15.29
7 7	13 13	13:30 13:45	0.513 0.612	1.766 1.712	12.790 15.359	22.59 26.29
7 7	13 13	14:00 14:15	0.615 0.582	1.748 1.736	15.437 14.577	26.98 25.31
7 7 7 7 7 7 7 7 7	13 13	14:30 14:45	0.539 0.493	1.738	13.461 12.275	23.40 21.86
7	13 13	15:00 15:15	0.451	1.665	11.198	18.64
7	13	15:30	0.398	1.460	10.406 9.847	15.19 6.71
7	13 13	15:45 16:00	0.373	0.438	9.213 8.783	4.04 1.05
7	13 13	19:30 19:45	0.221 0.209	0.154 0.331	5.403 5.105	0.83 1.69
7	13	20:00	0.203	0.307	4.957	1.52

Month 7	Date 13	Time (CDT) 20:15	Depth (ft) 0.202	Velocity (ft/sec) 0.232	Area (ft2) 4.932	Discharge (cfs) 1.14
7	13	20:30	0.199	0.065	4.858	0.32
7	13	20:45	0.190	0.333	4.635	1.54
7	13	21:00	0.187	0.312	4.561	1.42
7	13	21:15	0.190	0.052	4.635	0.24
7	13	21:30	0.183	0.058	4.462	0.26
7	13	21:45	0.175	0.018	4.265	0.08
7	14	11:45	0.229	0.703	5.601	3.94
7	14	12:00	0.483	1.461	12.018	17.56
7	14	12:15	0.525	1.409	13.100	18.46
7	14	12:30	0.518	1.323	12.919	17.09
7	14	12:45	0.497	1.334	12.378	16.51
7	14	13:00	0.465	1.167	11.556	13.49
7	14	13:15	0.439	0.969	10.891	10.55
7	14	13:30	0.413	0.969	10.228	9.91
7	14	13:45	0.400	0.574	9.898	5.68
, 7 7 7	14 14 14	14:00 14:15 14:30	0.386 0.369 0.363	0.412 0.321 0.046	9.542 9.112	3.93 2.92
7 7	14 14	14:45 15:00	0.347 0.329	0.073 0.117	8.960 8.556 8.102	0.41 0.62 0.95
7	14	15:15	0.319	0.069	7.850	0.54
7	14	15:30	0.305	0.076	7.499	0.57
7	14	15:45	0.297	0.076	7.298	0.55
7	14	16:00	0.284	0.084	6.973	0.59
7	14	16:15	0.276	0.073	6.773	0.49
7	14	16:30	0.268	0.047	6.573	0.31
7	14	16:45	0.263	0.036	6.448	0.23
7	14	17:00	0.256	0.019	6.273	0.12
7	17	17:00	0.203	0.760	4.957	3.77
7	17	17:15	0.256	0.868	6.273	5.45
7	17	17:30	0.262	0.844	6.423	5.42
7	17	17:45	0.264	0.932	6.473	6.03
7	17	18:00	0.465	1.537	11.556	17.76
7	17	18:15	0.675	1.894	17.010	32.22
7	17	18:30	0.706	1.959	17.828	34.92
7	17	18:45	0.673	1.888	16.958	32.02
7	17	19:00	0.620	1.768	15.568	27.52
7	17	19:15	0.567	1.683	14.187	23.88
7	17	19:30	0.519	1.573	12.945	20.36
7	17	19:45	0.478	1.553	11.890	18.46
7	17	20:00	0.442	1.497	10.968	16.42
7	17	20:15	0.410	1.468	10.152	14.90
7	17	20:30	0.381	1.371	9.415	12.91
7	17	20:45	0.354	1.338	8.732	11.68
7	17	21:00	0.331	1.163	8.152	9.48
7	17	21:15	0.311	1.035	7.649	7.92
7	17	21:30	0.294	0.839	7.223	6.06
7	17	21:45	0.274	0.886	6.723	5.96
7	17	22:00	0.259	0.723	6.348	4.59
7	17	22:15	0.255	0.571	6.248	3.57
7 7 7	17 17 17	22:30 22:45 23:00	0.243 0.236 0.231	0.399 0.286 0.159	5.949 5.775 5.651	2.37 1.65
, 7 7 7	18 18 18	17:45 18:00 18:15	0.193 0.385 0.488	0.837 1.320 1.685	4.709 9.517 12.146	0.90 3.94 12.56 20.47
7 7	18 18 18	18:30 18:45 19:00	0.816 0.967 0.991	2.337 2.422 2.331	20.755 24.838 25.494	48.50 60.16 59.43
7 7 7 7	18 18 18	19:15 19:30 19:45	0.943 0.855 0.759	2.252 2.090 1.964	24.184 21.802 19.233	54.46 45.57 37.77
7 7 7	18 18 18	20:00 20:15 20:30	0.668 0.584 0.518	1.847 1.768 1.666	16.826 14.629 12.919	31.08 25.86
7 7 7	18 18 18	20:30 20:45 21:00 21:15	0.462 0.414 0.376	1.577 1.469	11.479 10.254	21.52 18.10 15.06
7 7 7	18 18	21:30 21:45	0.342 0.312	1.400 1.315 1.199	9.289 8.430 7.675	13.00 11.08 9.20
7 7 7	18 18 18 18	22:00 22:15 22:30 22:45	0.292 0.267- 0.254 0.238	1.126 0.972 0.787	7.173 6.548 6.223	8.08 6.36 4.90
7	18	23:00	0.237	0.697 0.253	5.825 5.800	4.06 1.47

Month 7 7	Date 18 18	Time (CDT) 23:15 23:30	Depth (ft) 0.231 0.219	Velocit (ft/sec 0.088 0.098	_	Discharge (cfs) 0.50 0.52
7 7 7 7 7	18 31 31 31 31	23:45 16:00 16:15 16:30 16:45	0.205 0.906 2.517 3.403 3.644	0.112 0.540 2.561 3.702 1.842	5.006 23.179 71.126 101.168 109.790	0.56 12.52 182.15 374.52 202.23
7 7 7 7 7	31 31 31 31 31 31	17:00 17:15 17:30 17:45 18:00 18:15	3.468 2.937 2.521 2.123 1.773 1.493	4.435 2.818 3.937 3.879 3.283 3.367	103.474 85.042 71.256 58.604 47.912 39.652	458.91 239.65 280.53 227.32 157.30 133.51
7 7 7 7	31 31 31 31 31	18:30 18:45 19:00 19:15 19:30	1.263 1.089 0.927 0.817 0.727	2.932 2.856 2.323 2.435 2.379	33.061 28.192 23.749 20.781 18.384	96.93 80.52 55.17 50.60 43.73
7 7 7 7 7	31 31 31 31 31 31	19:45 20:00 20:15 20:30 20:45 21:00	0.648 0.578 0.526 0.468 0.431 0.391	2.794 2.471 2.392 2.302 2.285 2.070	16.301 14.473 13.125 11.633 10.687	45.54 35.76 31.40 26.78 24.42
7 7 7 7 7	31 31 31 31 31	21:15 21:30 21:45 22:00 22:15	0.354 0.324 0.299 0.278 0.251	1.769 1.975 1.870 2.019 1.816	9.669 8.732 7.976 7.348 6.823 6.149	20.01 15.45 15.75 13.74 13.77 11.17
7 7 7 8 8 8	31 31 31 1 1	22:30 22:45 23:00 10:30 10:45	0.234 0.217 0.202 0.167 0.399	1.920 1.708 1.599 1.834 2.325	5.726 5.304 4.932 4.068 9.872	10.99 9.06 7.89 7.46 22.95
8 8 8 8	1 1 1 1	11:00 11:15 11:30 11:45 12:00 12:15	0.529 0.674 0.569 0.533 0.582 0.691	2.262 2.243 1.746 1.861 2.392 2.137	13.203 16.984 14.239 13.306 14.577 17.432	29.86 38.10 24.86 24.76 34.87 37.25
8 8 8 8	1 1 1 1	12:30 12:45 13:00 13:15 13:30	0.796 2.675 4.254 4.861 4.471	2.977 3.181 3.318 3.589 4.146	20.219 76.292 132.477 156.278 140.845	60.19 242.69 439.56 560.88 583.94
8 8 8 8 8	1 1 1 1 1	13:45 14:00 14:15 14:30 14:45 15:00	3.825 3.179 2.648 2.223 1.873 1.600	3.785 3.776 4.332 4.023 3.416 4.018	116.393 93.326 75.404 61.733 50.925 42.778	440.55 352.40 326.65 248.35 173.96 171.88
8 8 8 8	1 1 1 1	15:15 15:30 15:45 16:00 16:15	1.362 1.180 1.041 0.910 0.816	3.368 3.348 3.157 3.127 3.080	35.876 30.726 26.866 23.287 20.755	120.83 102.87 84.82 72.82 63.92
8 8 8 8 8	1 1 1 1 1	16:30 16:45 17:00 17:15 17:30 17:45	0.727 0.646 0.582 0.533 0.489 0.446	2.775 2.630 2.358 2.569 2.440 2.386	18.384 16.248 14.577 13.306 12.172 11.070	51.01 42.73 34.37 34.18 29.70 26.41
8 8 8 8	1 1 2 2 2	18:00 18:15 13:45 14:00 14:15	0.410 0.368 1.270 1.994 2.225	2.130 2.165 3.474 3.708 3.848	10.152 9.086 33.259 54.616 61.796	21.62 19.67 115.54 202.52 237.79
8 8 8 8 8	2 2 2 2 2 2	14:30 14:45 15:00 15:15 15:30	2.305 2.327 2.290 2.273 2.217	3.862 4.088 4.172 3.934 4.256	64.324 65.023 63.848 63.310 61.544	248.42 265.81 266.38 249.06 261.93
8 8 8	2 2 2	15:45 16:00 16:15 16:30	2.062 1.870 1.637 1.428	3.897 4.224 3.631 3.720	56.711 50.834 43.867 37.771	221.00 214.72 159.28 140.51

Month	Date	Time	Depth (ft)	Velocity		Discharge
8	2	(CDT) 16:45	1.246	(ft/sec) 3.421	32.581	(cfs) 111.46
8	2	17:00	1.073	3.454	27.749	95.85
8	2	17:15	1.169	3.234	30.418	98.37
8	2	17:30	1.281	3.192	33.570	107.16
8	2	17:45	1.372	3.767	36.163	
8	2	18:00	1.446	3.690	38.291	136.22 141.29
8	2	18:15	1.438	3.762	38.060	143.18
8		18:30	1.353	3.715	35.619	132.32
8	2 2	18:45	1.225	3.274	31.989	104.73
8	2	19:00	1.104	3.308	28.608	94.63
8		19:15	0.976	3.153	25.083	79.09
8	2	19:30	0.856	2.935	21.829	64.07
8	2	19:45	0.757	2.851	19.180	54.68
8 8	2	20:00 20:15	0.684 0.609	2.677 2.631	17.247 15.280	46.17
8	2	20:30	0.552	2.574	13.798	40.20 35.52
8 8	2 2 2	20:45 21:00	0.507 0.461	2.360 2.444	12.635 11.454	29.82 27.99
8	2	21:15	0.433	2.433	10.738	26.13
8	2	21:30	0.394	2.330	9.745	22.71
8	2	21:45	0.376	2.013	9.289	18.70
8	2	22:00	0.345	2.194	8.505	18.66
8		22:15	0.316	2.118	7.775	16.47
8	2 2 2	22:30 22:45	0.302 0.293	2.100 2.093	7.424 7.198	15.59 15.07
8	2	23:00	0.286	1.935	7.023	13.59
8	2	23:15	0.268	1.909	6.773	12.93
8	2	23:30		1.986	6.573	13.05
8	2	23:45	0.256	2.132	6.273	13.37
8	3	0:00	0.264	2.035	6.473	13.17
8	3	0:15	0.263	1.957	6.448	12.62
8		0:30	0.273	2.007	6.698	13.44
8	3	0:45	0.282	1.993	6.923	13.80
8	3	1:00	0.270	2.124	6.623	14.07
8		1:15	0.259	2.015	6.348	12.79
8	3	1:30	0.233	1.859	5.701	10.60
8	3	1:45	0.220	1.574	5.378	8.46
8	6	11:30 11:45	0.167 0.784	1.235 1.385	4.068 19.899	5.02
8	6	12:00	1.121	2.478	29.080	27.56 72.06
8	6	12:15	1.166	2.666	30.334	80.87
8	6	12:30	1.266	3.131	33.146	103.78
8	6	12:45	1.218	2.728	31.792	86.73
	6	13:00	1.151	2.927	29.915	87.56
8	6	13:15 13:30	1.203	2.984	31.371	93.61
8	6	13:45	1.143	3.056 3.066	31.792 29.692	97.16 91.04
8	6	14:00	1.028	2.841	26.508	75.31
8	6	14:15	0.896	2.792	22.908	63.96
8	6	14:30	0.770	2.485	19.526	48.52
8	6	14:45	0.673	2.558	16.958	43.38
8	6	15:00	0.599	2.237	15.020	33.60
8	6	15:15	0.518	2.136	12.919	27.59
8	6	15:30	0.462	2.161	11.479	24.81
8	6	15:45	0.415	2.116	10.279	21.75
8	6	16:00	0.367	2.052	9.061	18.59
8	6	16:15	0.333	1.987	8.203	16.30
8	6	16:30 16:45	0.306 0.276	2.025 1.912	7.524 6.773	15.24 12.95
8	б	17:00	0.255	1.768	6.248	11.05
8	6	17:15	0.234	1.879	5.726	10.76
8	6	17:30	0.214	1.831	5.229	9.57
8	6	17:45	0.199	1.743	4.858	8.47
8	7	14:30	0.415	1.807	10.279	18.57
8	7	14:45	0.445	1.963	11.044	21.68
8	7	15:00	0.349	1.763	8.606	15.17
8	7	15:15	0.275	1.821	6.748	12.29
8	7	15:30	0.335	2.121	8.253	17.50
8 8	7 7	15:45 16:00	0.407	2.217	10.076	22.34
8	7	16:15	0.519	2.216 2.320	12.945 14.187	28.69 32.91
8	7	16:30	0.556	2.213	13.902	30.76
8	7	16:45	0.518	2.119	12.919	27.38
8	7	17:00	0.470	2.084	11.684	24.35
8	7	17:15	0.428	2.086	10.610	22.13
8	7	17:30	0.377	2.081	9.314	19.38

Month	Date	Time	Depth (ft)	Velocit (ft/sec		Discharge (cfs)
8	7 7	17:45 18:00	0.342	1.929	8.430 7.574	16.26 14.25
8 8	7 7	18:15 18:30	0.281 0.250	1.940 1.710	6.898 6.124	13.38 10.47
8	7 7	18:45 19:00	0.228 0.209	1.818 1.821	5.577 5.105	10.14 9.30
8	7	19:15 19:30	0.208 0.199	0.000	5.080 4.858	0.00
8	9	14:30 14:45	0.223 0.752	1.794 2.517	5.452 19.047	9.78 47.94
8	9	15:00 15:15	0.940 0.922	2.591 2.298	24.102 23.613	62.45 54.26
8	9	15:30 15:45	0.832	2.213 2.206	21.184 18.569	46.88 40.96
8 8 8	9 9 9	16:00 16:15	0.639	2.091 1.977	16.065 13.824	33.59 27.33
8 8	9	16:30 16:45 17:00	0.500 0.446 0.411	2.174 1.986	12.455 11.070	27.08 21.98
8	9 9	17:15 17:30	0.368	1.876 1.803 1.873	10.177 9.086 8.404	19.09 16.38
8	9 9	17:45 18:00	0.314	1.913	7.725 7.073	15.74 14.78 12.34
8 8	9 9	18:15 18:30	0.264	1.736	6.473 5.900	11.24
8 8	9 9	18:45 19:00	0.227	1.714	5.552 5.056	9.52 6.61
8 8	9 9	19:15 19:30	0.195 0.167	1.437 1.377	4.759	6.84 5.60
8	9 10	19:45 14:15	0.190 0.419	0.000 1.721	4.635 10.381	0.00 17.87
8	10 10	14:30 14:45	$\frac{1.145}{1.507}$	2.515 2.620	29.748 40.058	74.82 104.95
8 8	10 10	15:00 15:15	1.497	2.898 2.732	39.768 36.191	115.25 98.87
8 8 8	10 10	15:30 15:45	1.223	2.701 2.525	31.933 27.280	86.25 68.88
8 8	10 10 10	16:00 16:15	0.924	2.436	23.667	57.65 50.55
8	10 10	16:30 16:45 17:00	0.691 2.733	2.479 1.919	17.432 78.210	43.21 150.08
8	10 10	17:15 17:30	3.509 3.439 3.029	2.879 3.621 3.035	104.936	302.11 370.95
8	10 10	17:45 18:00	2.738	3.343 3.728	88.169 78.375 67.386	267.59 262.01 251.21
8	10 10	18:15 18:30	2.068	2.951 3.536	56.897 48.572	167.90 171.75
8	10 10	18:45 19:00	1.583	3.038 3.313	42.278 36.793	128.44 121.90
8 8	10 10	19:15 19:30	1.231 1.089	2.967 2.855	32.158 28.192	95.41 80.49
8 8	10 10	19:45 20:00	0.960 0.861	2.661 2.954	24.647 21.963	65.58 64.88
8	10 10	20:15 20:30	0.766 0.674	2.829 2.635	19.419 16.984	54.94 44.75
8	10 10	20:45 21:00	0.596 0.541	2.426 2.486	14.941 13.513	36.25 33.59
8 8 8	10 10	21:15 21:30	0.492 0.450	2.331 2.289	12.249 11.172	28.55 25.57
8	10 10	21:45 22:00	0.414	2.102 2.189	10.254 9.339	21.55 20.44
8 8 8	10 10	22:15	0.352	2.027	8.682 8.102	17.60 16.20
8	10 10 10	22:45 23:00 23:15	0.310	2.011	7.624 6.748	15.33 12.54
8	10 11	23:30 9:45	0.254 0.242 0.597	1.655	6.223 5.925	10.30
8 8	11 11	10:00 10:15	0.959 1.571	1.800 2.558 2.864	14.967 24.619	26.94 62.98
8	11 11	10:30 10:45	1.899	3.266 3.570	41.927 51.714 59.040	120.08 168.90
8	11 11	11:00 11:15	2.220	3.827 3.367	61.639 59.102	210.77 235.89 199.00
8	11	11:30	2.024	3.841	55.538	213.32

Month	Date	Time (CDT)	Depth (ft)	Velocity (ft/sec)	Area (ft2)	Discharge (cfs)
8	11	11:45	1.856	3.547	50.411	178.81
8	11	12:00	1.685	3.343	45.288	151.40
8	11	12:15	1.527	3.112	40.641	126.47
8 8	11	12:30	1.354	3.121	35.648	111.26
8	11	12:45	1.220	3.190	31.848	101.60
	11	13:00	1.082	3.032	27.998	84.89
8	11	13:15	0.955	2.959	24.510	72.53
8	11	13:30	0.858	2.770	21.883	60.62
8	11	13:45	0.770	2.808	19.526	54.83
8	11	14:00	0.697	2.626	17.590	46.19
8	11	14:15	0.629	2.424	15.803	38.31
8	11	14:30	0.580	2.570	14.525	37.33
8	11	14:45	0.530	2.391	13.229	31.63
8	11	15:00	0.488	2.393	12.146	29.07
8	11	15:15	0.453	2.230	11.249	25.09
8	11	15:30	0.416	2.251	10.305	23.20
8	11	15:45	0.391	2.092	9.669	20.23
8	11	16:00	0.358	2.089	8.833	18.45
8	11	16:15	0.343	2.005	8.455	16.95
8	11	16:30	0.319	1.950	7.850	15.31
8	11	16:45	0.294	1.841	7.223	13.30
8	11 11	17:00	0.282	1.988	6.923	13.76
8	11	17:15 17:30	0.256 0.247	1.832 1.802	6.273 6.049	11.49 10.90
8	11	17:45	0.231	1.686	5.651	9.53
8	11	18:00	0.218	1.783	5.328	9.50
8	12	6:45	0.265	1.769	6.498	11.49
8	12	7:00	0.482	2.292	11.992	27.49
8	12	7:15	0.531	2.428	13.255	32.18
8	12	7:30	0.607	2.518	15.228	38.34
8	12	7:45	0.678	2.581	17.089	44.11
8	12	8:00	0.736	2.684	18.622	49.98
8	12	8:15	0.765	2.790	19.393	54.11
8	12	8:30	0.755	2.411	19.127	46.11
8	12	8:45	0.739	2.566	18.702	47.99
8	12	9:00	0.712	2.295	17.986	41.28
8	12	9:15	0.657	2.380	16.537	39.36
· 8	12	9:30	0.603	2.535	15.124	38.34
	12	9:45	0.554	2.372	13.850	32.85
8	12	10:00	0.506	2.319	12.609	29.24
8	12	10:15	0.463	2.219	11.505	25.53
8	12	10:30	0.426	2.265	10.559	23.92
8	12	10:45	0.385	2.042	9.517	19.43
8	12	11:00	0.352	1.996	8.682	17.33
8 8	12 12	11:15 11:30	0.327 0.304	1.885	8.052	15.18
8	12	11:45	0.284	1.924 1.814	7.474 6.973	14.38 12.65
8	12	12:00	0.261	1.797	6.398	11.50
8	12	12:15	0.257	1.876	6.298	11.82
8	12	12:30	0.251	1.774	6.149	10.91
8	12	12:45	0.242	1.817	5.925	10.76
8	12	13:00	0.226	1.663	5.527	9.19
8	12	13:15	0.216	1.581	5.279	8.35
8	12	13:30	0.214	1.819	5.229	9.51
8	12	13:45	0.201	1.548	4.907	7.60
8	12	14:00	0.197	1.615	4.808	7.77
8	12	18:45	0.188	1.621	4.586	7.43
8	12	19:00	0.337	1.881	8.303	15.62
	12	19:15	0.359	1.815	8.859	16.08
8	12	19:30	0.336	1.878	8.278	15.55
8	12	19:45	0.322	1.817	7.926	14.40
8	12	20:00	0.289	1.814	7.098	12.88
8	12	20:15	0.258	1.630	6.323	10.31
8	12	20:30	0.239	1.587	5.850	9.28
8	12	20:45	0.221	1.657	5.403	8.95
8	12	21:00	0.203	1.687	4.957	8.36

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